



Influence of groundwater residence time on biogeochemical transformations after legacy sediment removal from a headwater stream in Lancaster County, Pennsylvania

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Research problem

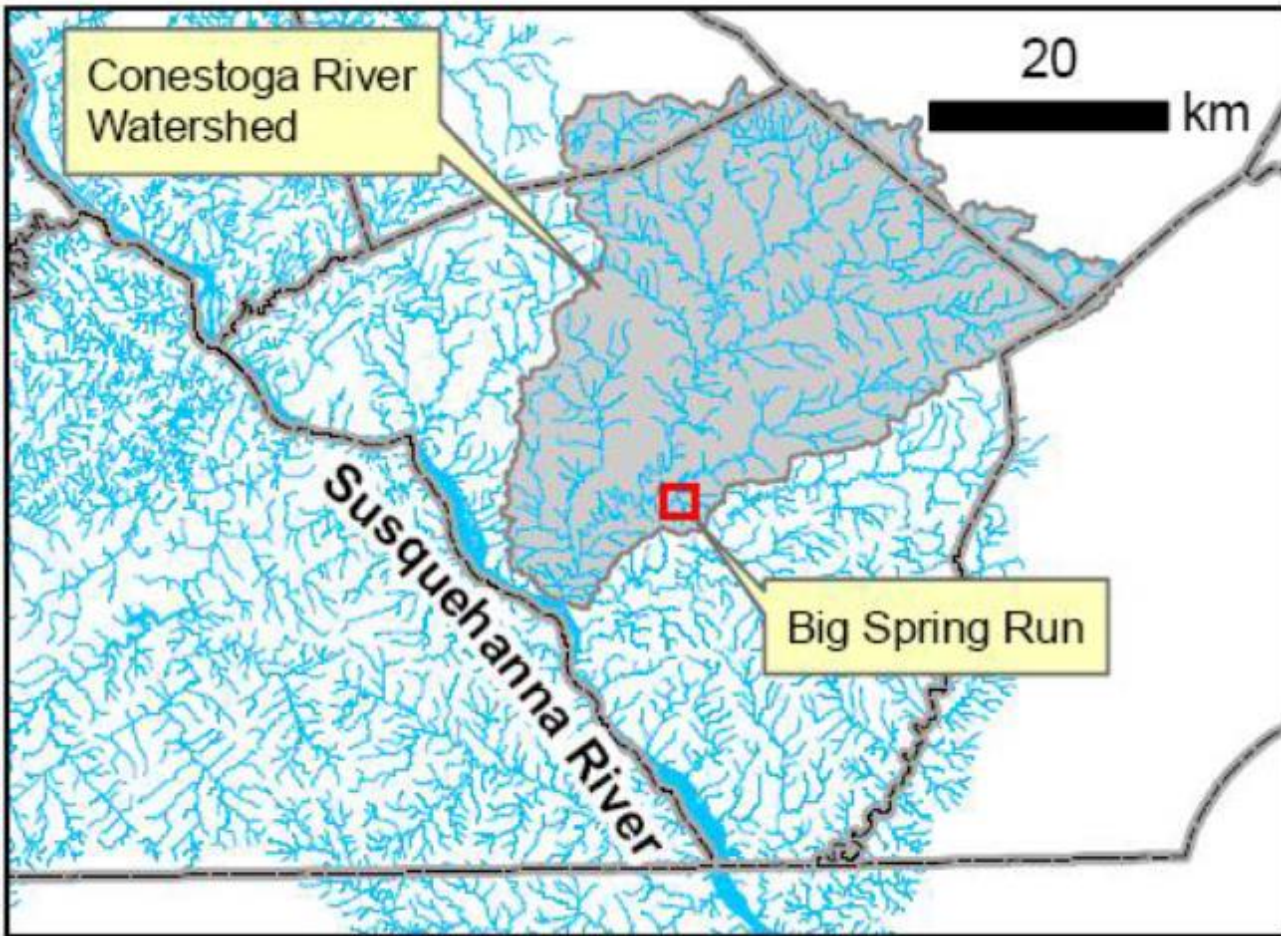
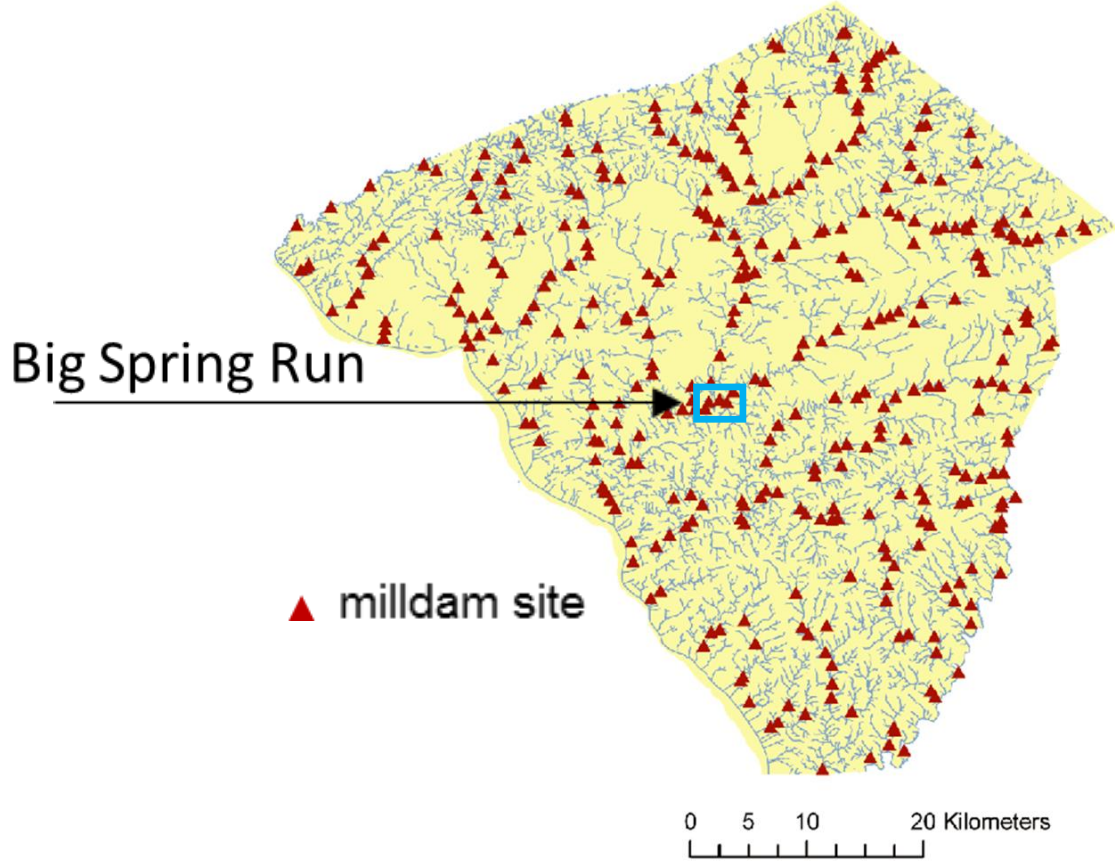


- Mounting evidence suggests that concurrent, widespread deforestation and milldam construction between the 18th and 20th centuries blanketed mid-Atlantic Piedmont valley bottoms with a thick (~1.5 m - >10 m) accumulation of fine-grained legacy sediment.
- Banks of numerous 1st- to 3rd-order streams became incised as dams were breached. Consequentially, hydrological and biogeochemical processes between the former floodplain and active stream channel were decoupled. Given the ubiquity of milldam-impacted streams in the Chesapeake Bay, it is critical to understand:

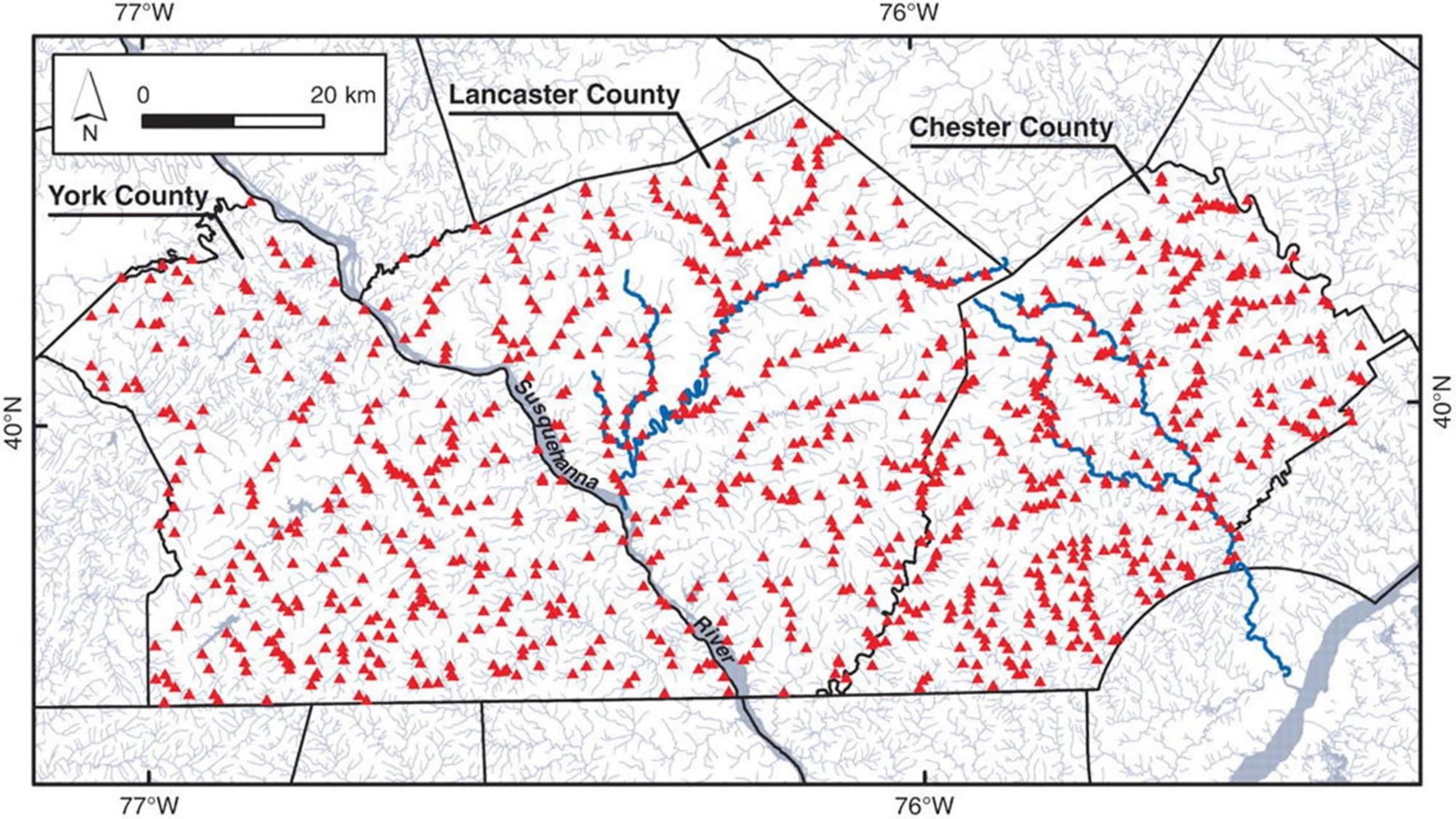
How does a restored hydrologic connection in this milldam-impacted landscape influence biogeochemical processing of nitrogen and carbon?

- We present findings from an 8-year, pre- and post-restoration study conducted at Big Spring Run, Lancaster, Pennsylvania where we evaluated variability in the oxygen isotope composition of catchment waters to examine how groundwater contact times and flowpaths influence seasonal nitrate ($\text{NO}_3\text{-NO}_2\text{-N}$) and dissolved organic carbon (DOC) dynamics.

LOCATION OF 383 HISTORIC MILL DAMS, LANCASTER COUNTY, PA



Equals 1 mill every 2.4 km of stream length in County



Big Spring Run - Wetland Restoration



Bank erosion in the upper reaches of BSR, made worse during high water events (Wolman 1959; Merritts *et al.* 2010) and freeze-thaw cycles (Merritts *et al.* 2011; Inamdar *et al.* 2017; Walter *et al.* 2018), resulted in an **average annual release of 94 tons of fine suspended sediment between 2008 and 2011** (PA DEP 2013).

Typical incised channel, erosional banks

9/13/2011



PA DEP began discussion with landowners on two adjacent farms to designate a portion of the BSR catchment that would eventually become the site of an 8-year study to pilot a BMP designed to restore streams in the eastern US impacted by historic milldams to a close approximation of their natural stream morphology (Hartranft *et al.* 2011).

Natural valley morphology

07/27/2012

The restoration

- LandStudies, a local engineering firm based in Lancaster, began the restoration in October 2011
- 21,721 cubic feet of legacy sediment was removed from a 4 km² portion of the catchment
- 1,479 linear feet of stream valley and 3,074 linear feet of stream channel was [restored](#)

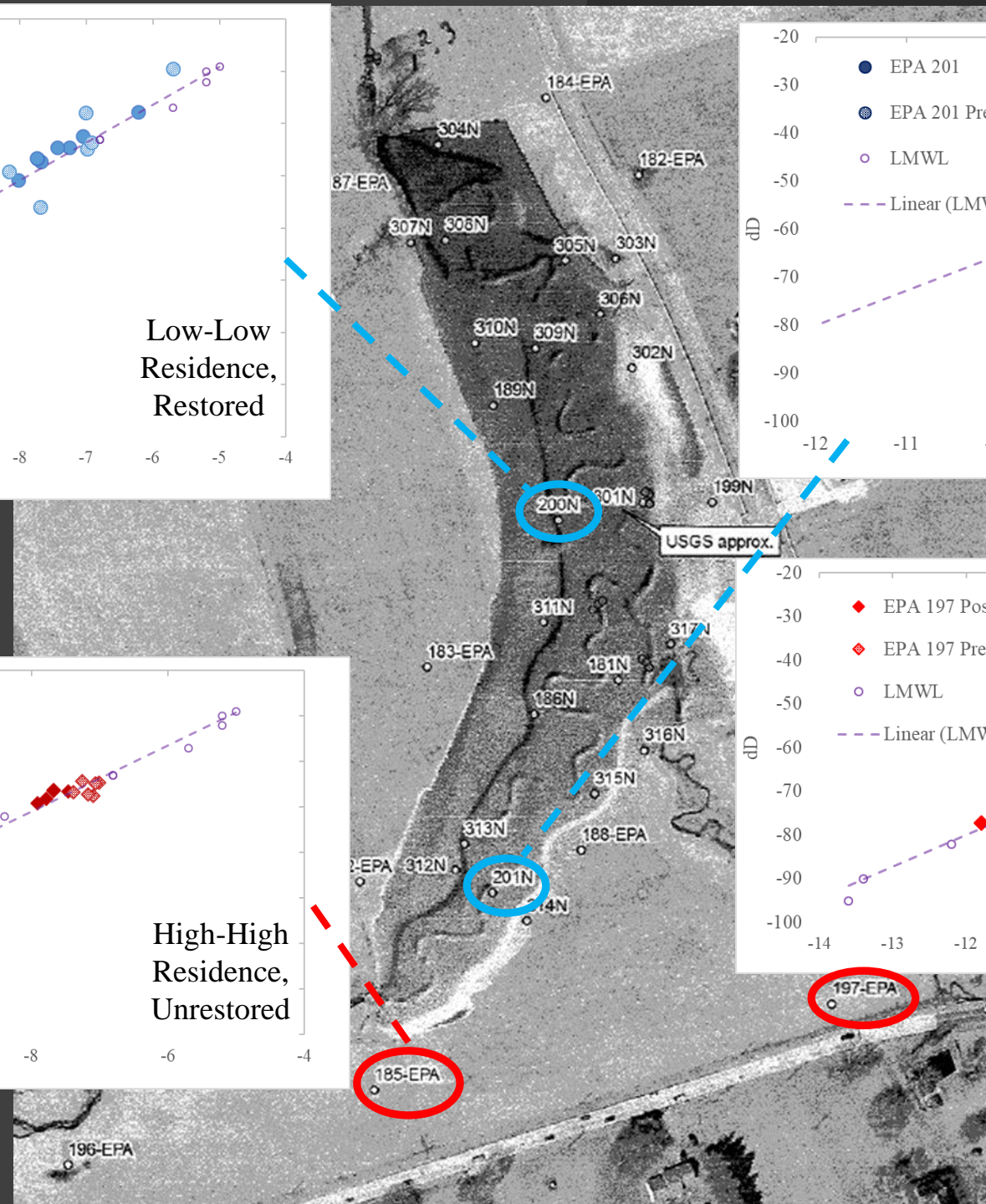
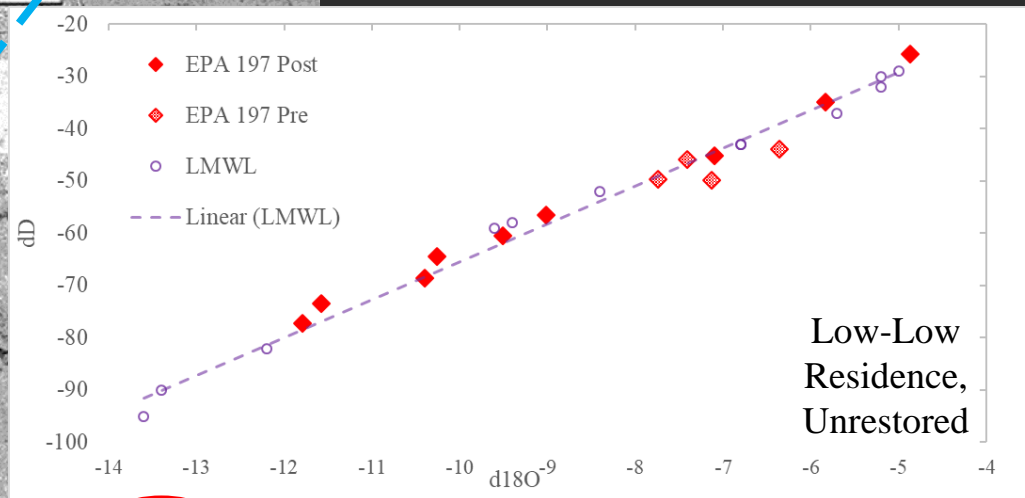
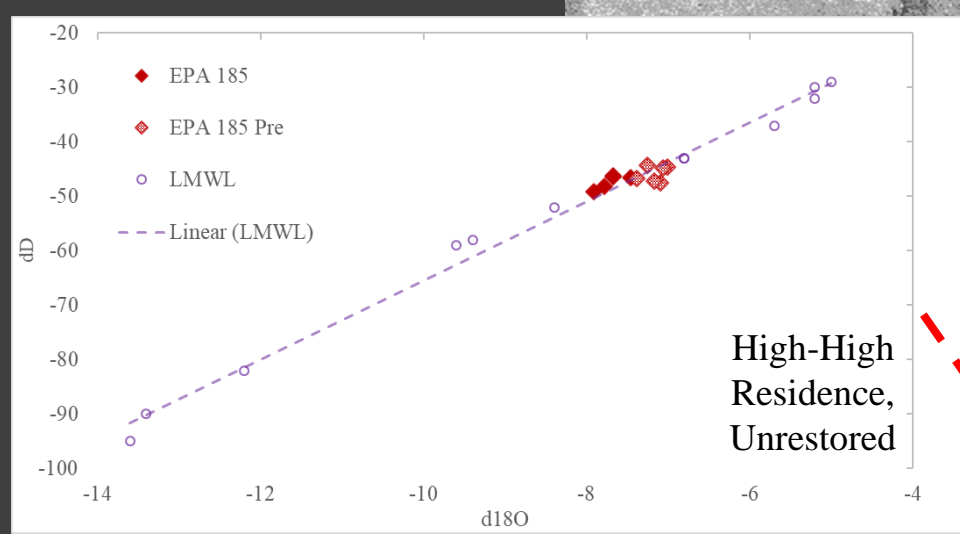
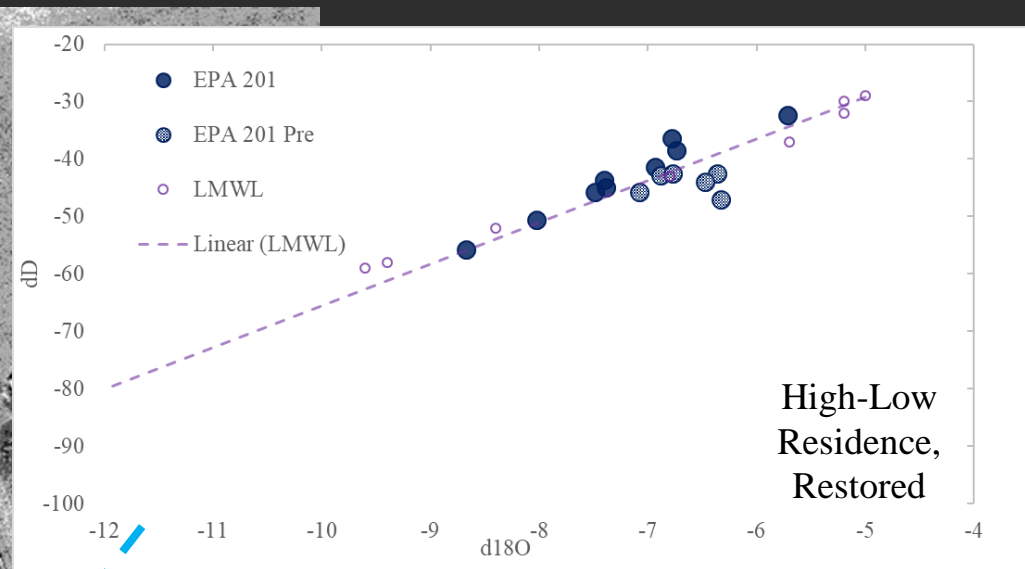
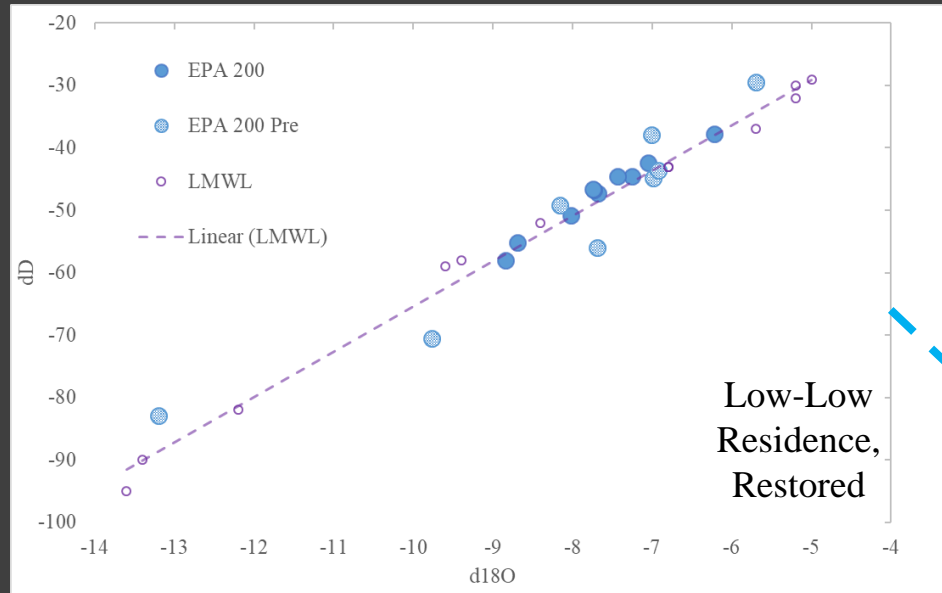


Isotope theory

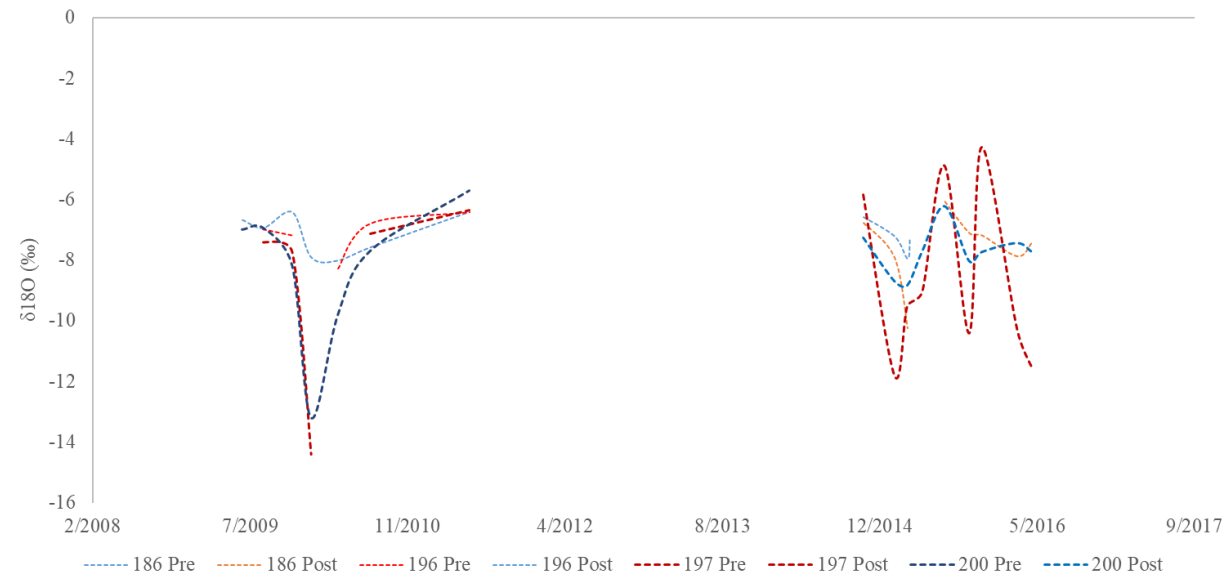
- We utilized the variance in the stable isotope ratio (¹⁸O/¹⁶O) in water to determine a relative ranking of groundwater residence time for 13 wells in the catchment pre- and post-restoration

$$\delta^{18}\text{O} = \left(\frac{\left(\frac{^{18}\text{O}}{^{16}\text{O}} \right)_{\text{sample}}}{\left(\frac{^{18}\text{O}}{^{16}\text{O}} \right)_{\text{standard}}} - 1 \right) \times 1000 \text{ ‰}$$

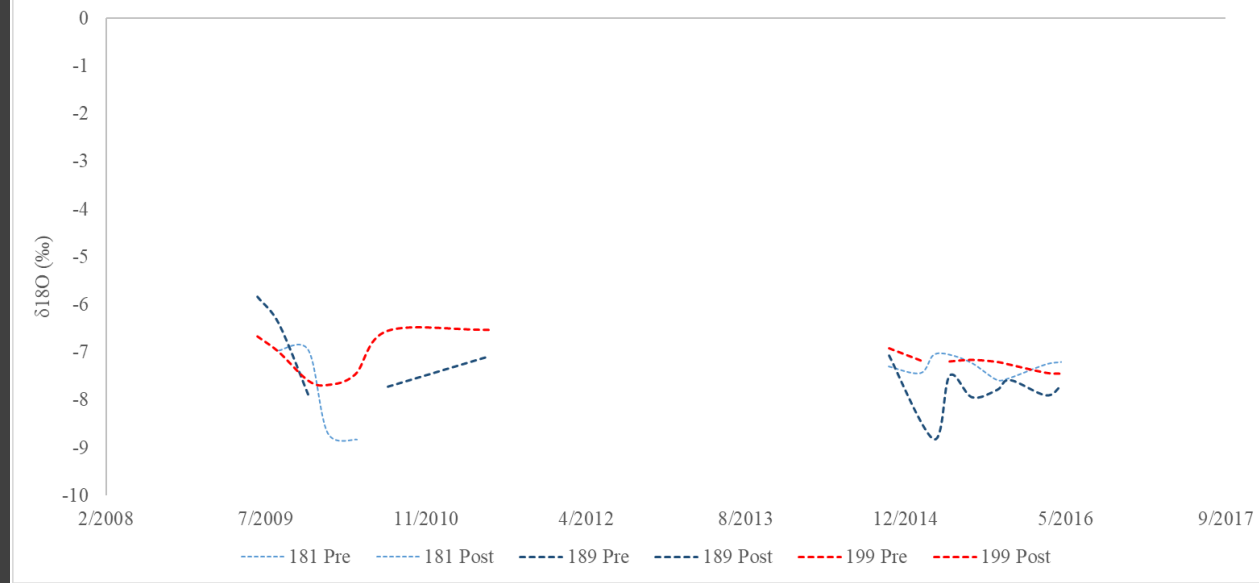
- Water samples were taken on a bi-monthly basis for 2 years pre-restoration (2009-2011) and post-restoration (2014-2016)
- Greater variance in $\delta^{18}\text{O}$ infers shorter groundwater contact times and a more active flow path



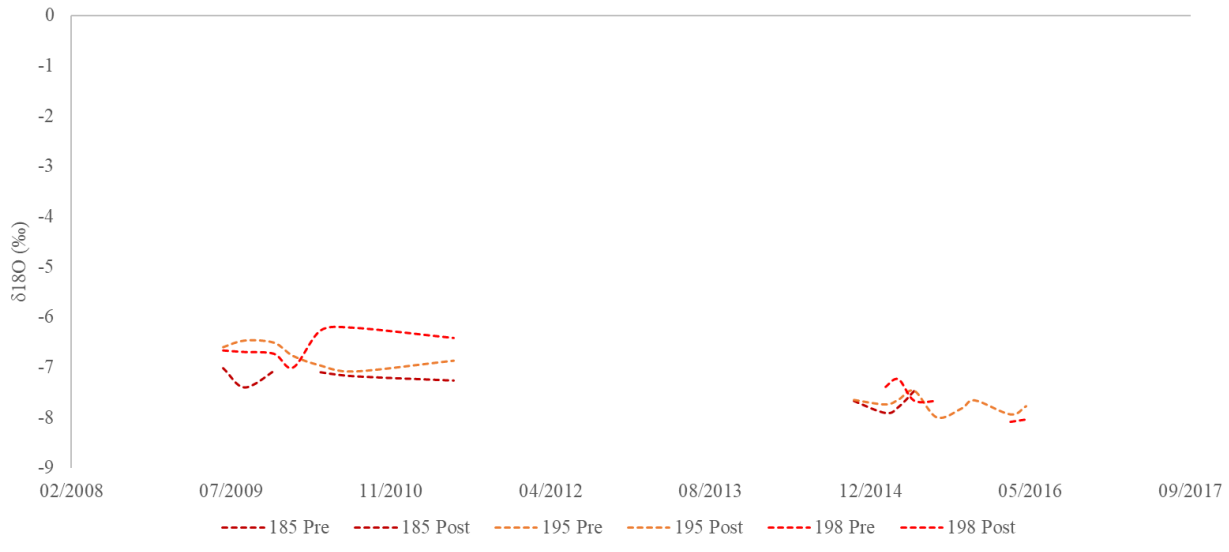
Low - Low Residence, $\delta^{18}\text{O}$ (‰)



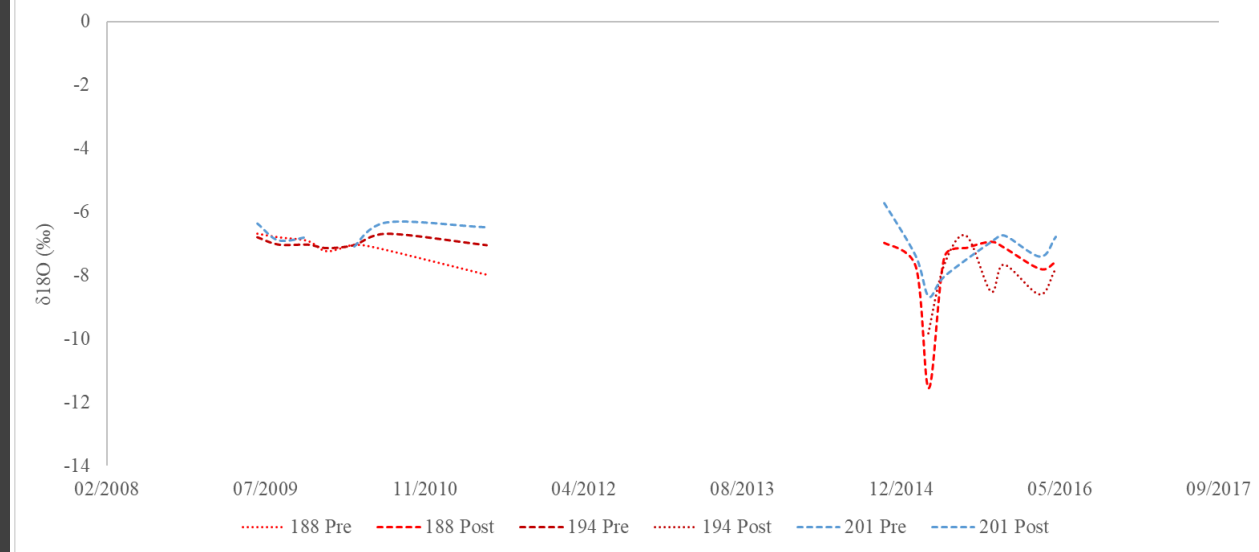
Low - High Residence, $\delta^{18}\text{O}$ (‰)



High - High Residence, $\delta^{18}\text{O}$ (‰)



High - Low Residence, $\delta^{18}\text{O}$ (‰)



Introduction: Low-Low Residence

Groundwater residence time at these wells remained *faster* than the catchment average pre- and post-restoration.

Groundwater at these locations assumes relatively *active flow paths*.

Wells **186** and **200** are both located inside the **restored area** whereas **196** and **197** are located **outside the restored area**.

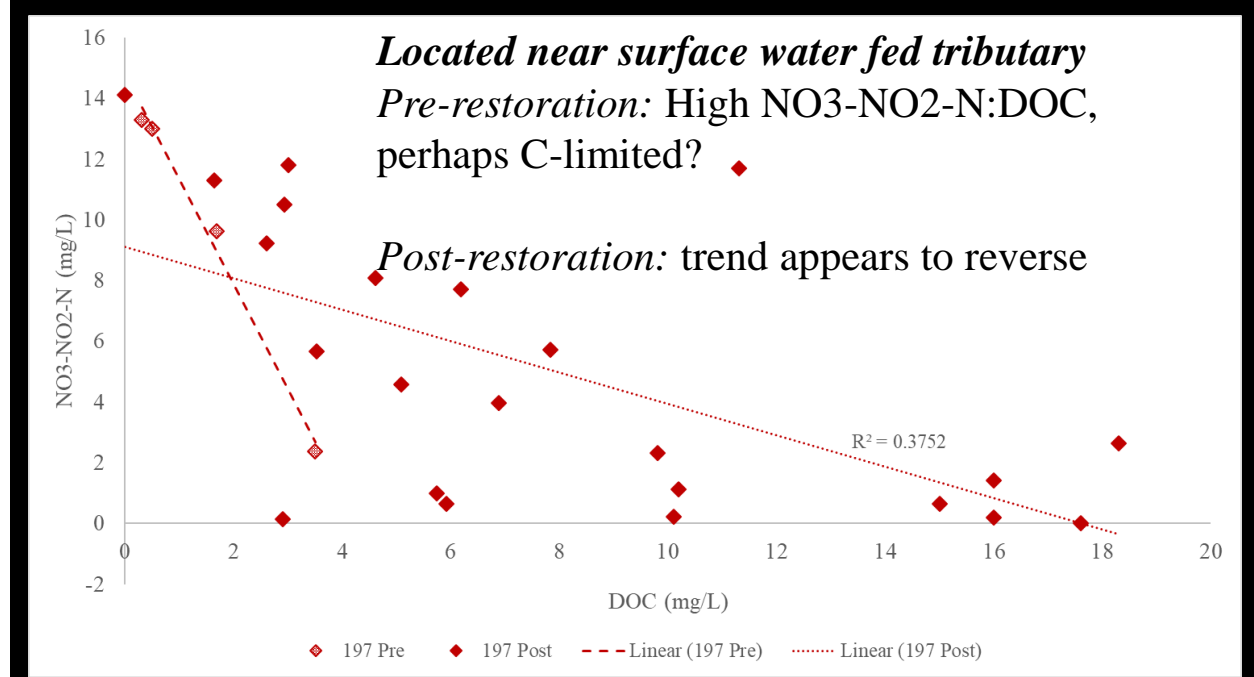
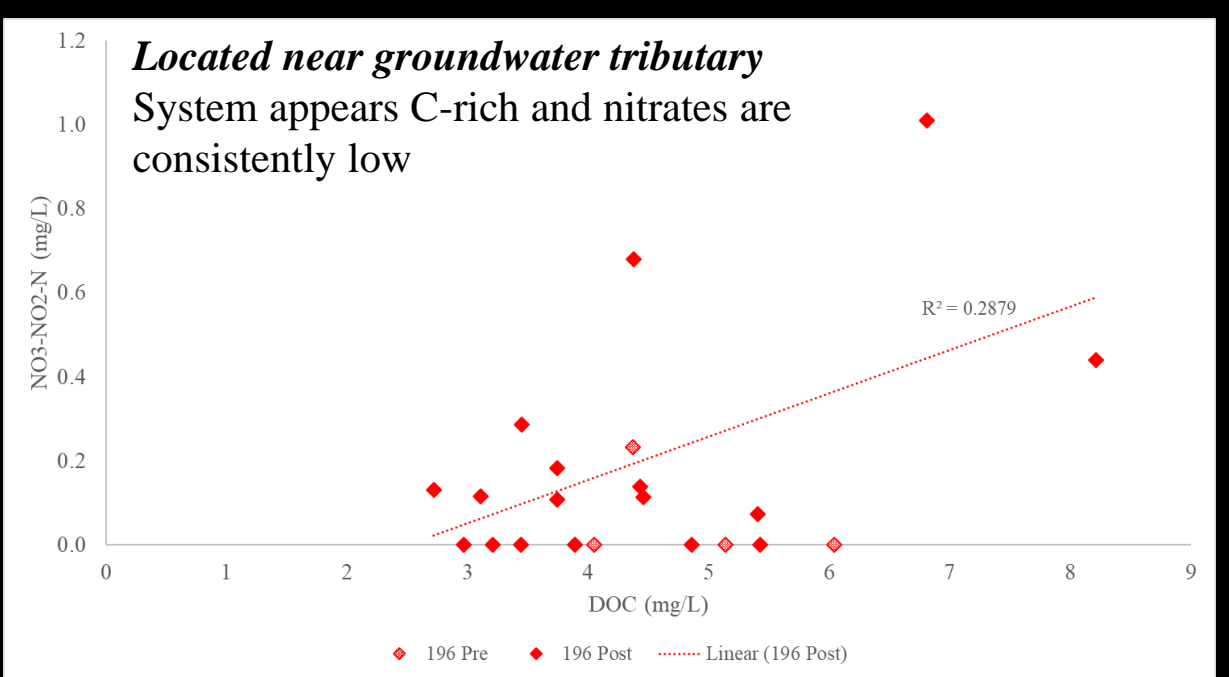
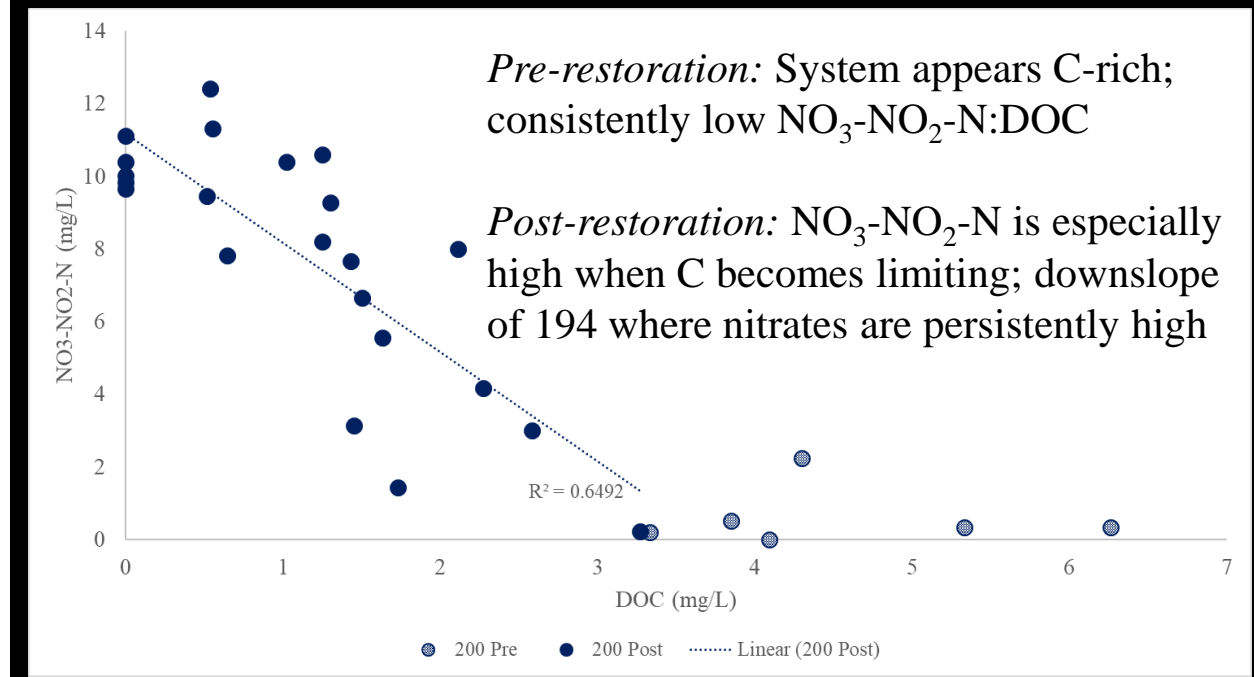
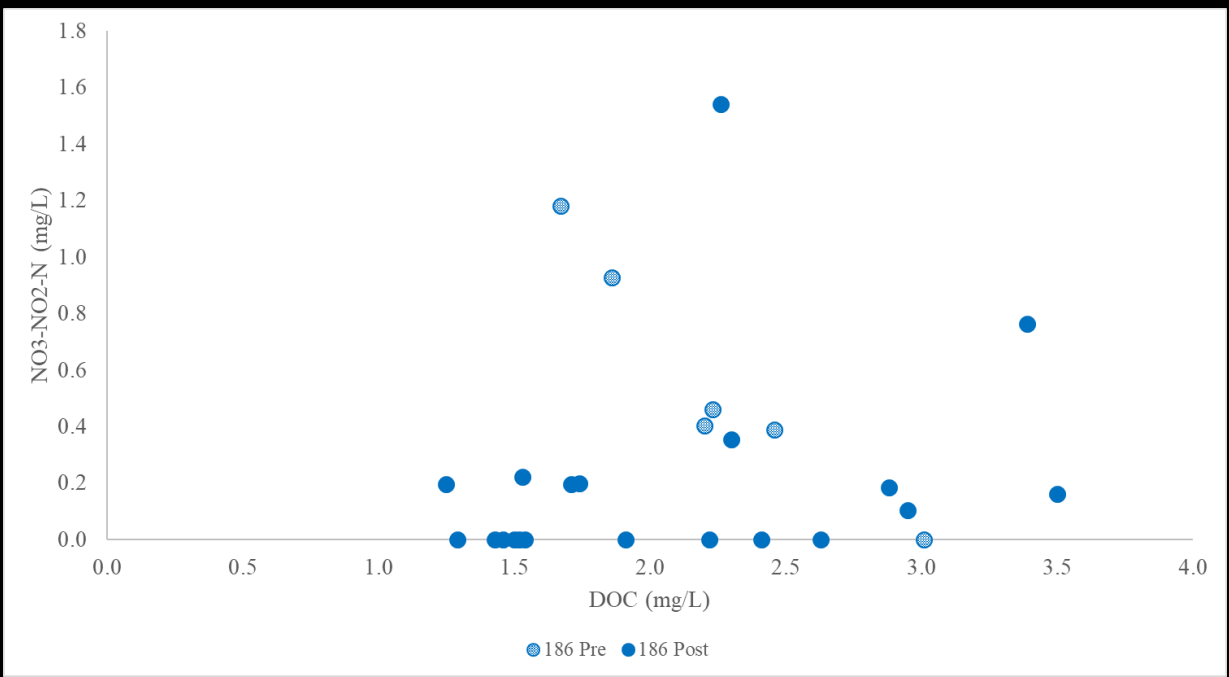
Well **186** had a depth of 7.2 ft, but was replaced post-restoration so the depth became 3.5 ft. The screened interval (SI) begins at 2.5 ft (relict) - The soil type is Newark.

Well **200** had a pre-restoration depth of 3.5 ft but was re-drilled to 4.0 ft post-restoration. The SI begins at 3 ft (relict). The soil type is Newark.

Well **197** may be seasonally influenced by road salts and/or leaf litter decomposition in stormwater runoff due to its roadside proximity. It is also located closer to trees than many other wells in the catchment. This well has a depth of 18.8 ft. The SI begins at 13.5 ft (bedrock). The overlying soil type is Newark.

Well **196** has a depth of 8.3 ft (bedrock), SI at 5 ft and is located in close proximity to the groundwater-fed tributary. The overlying soil type is Newark.





Introduction: Low-High Residence

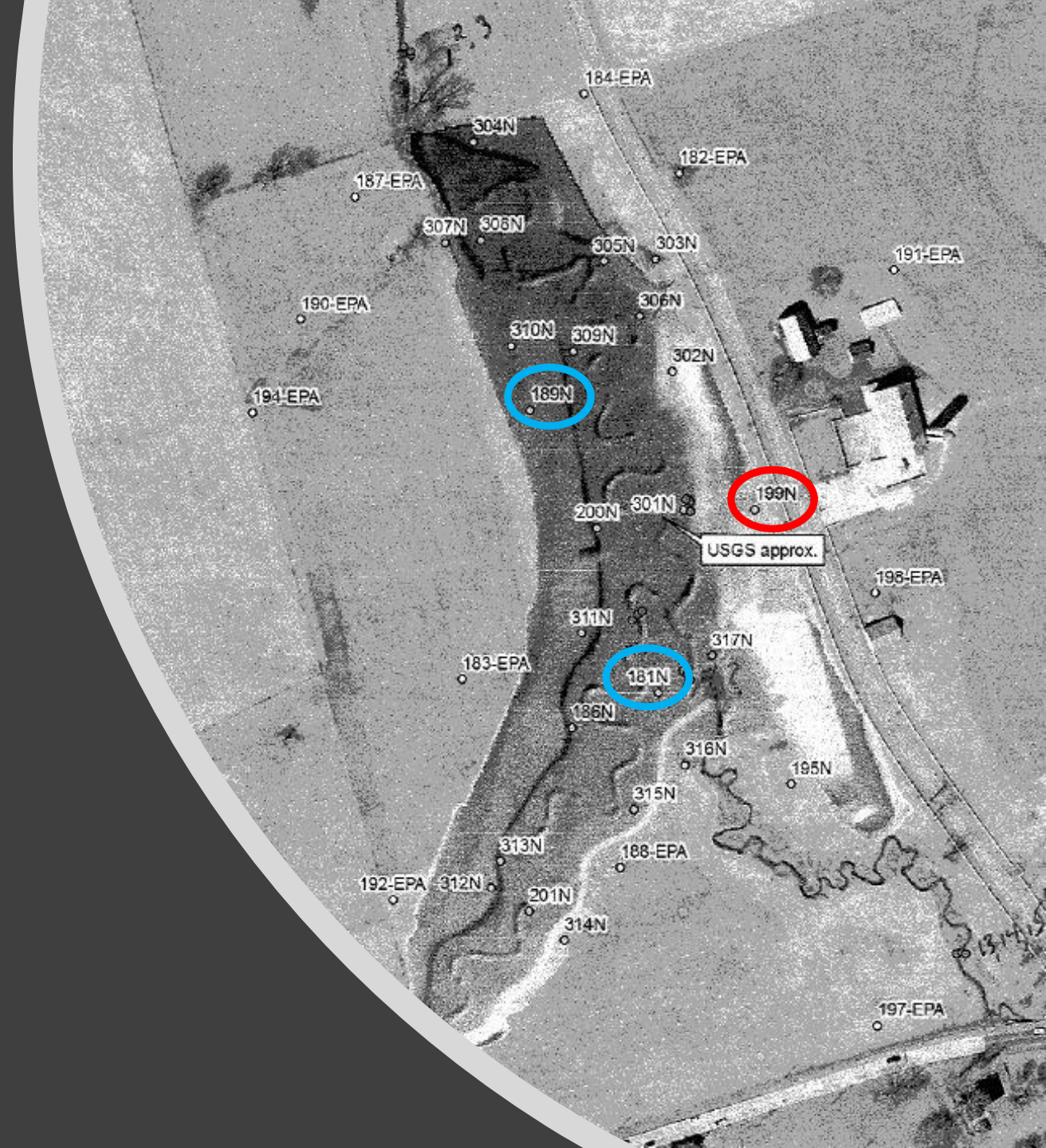
Groundwater residence time at these wells was faster than the catchment average pre-restoration, but slower than the catchment average after the restoration occurred. Thus, the **groundwater flow paths at these wells became *less active* post-restoration.**

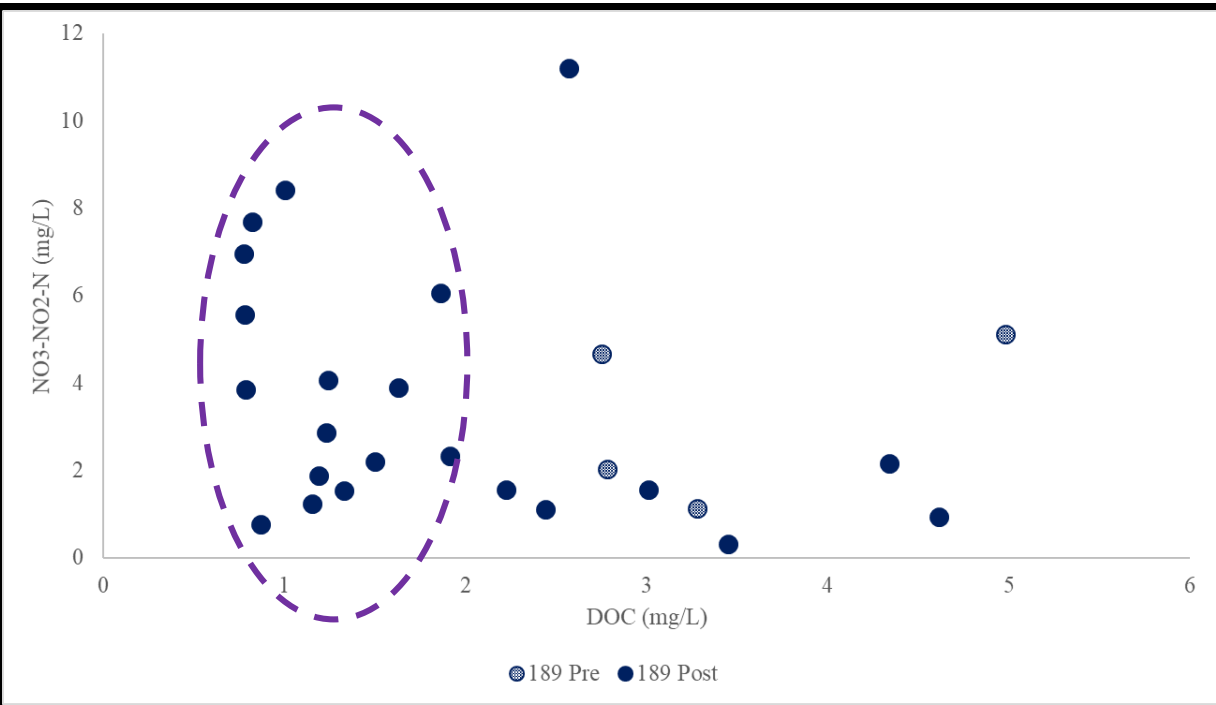
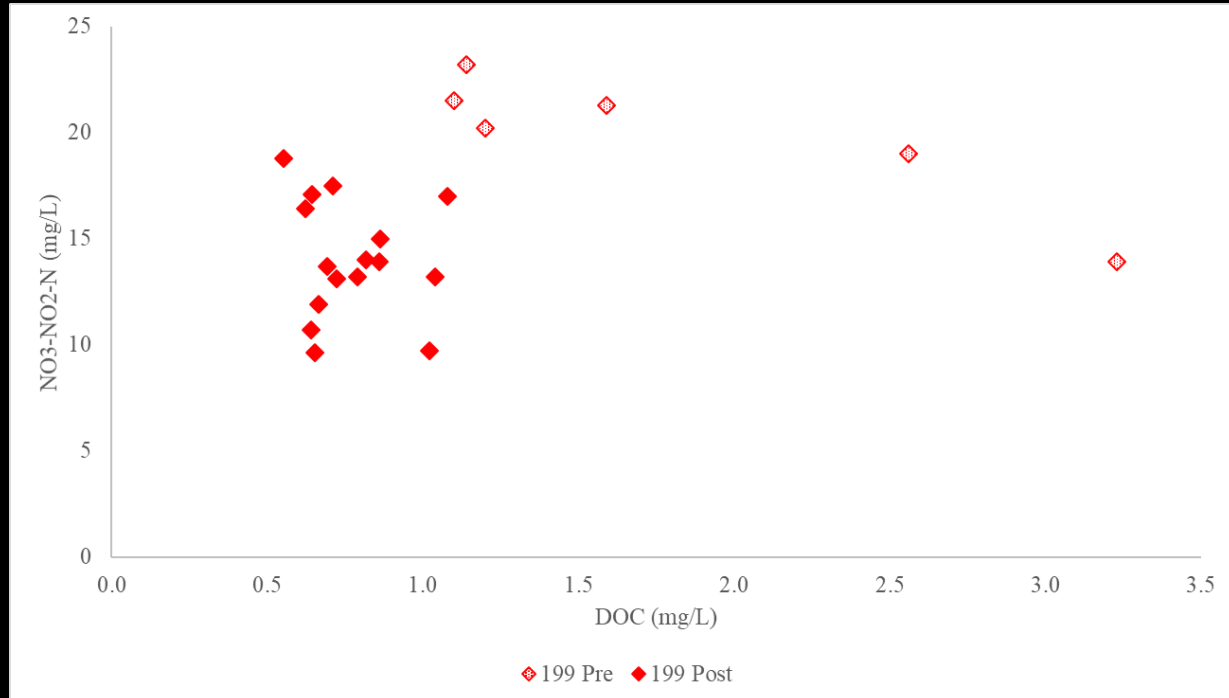
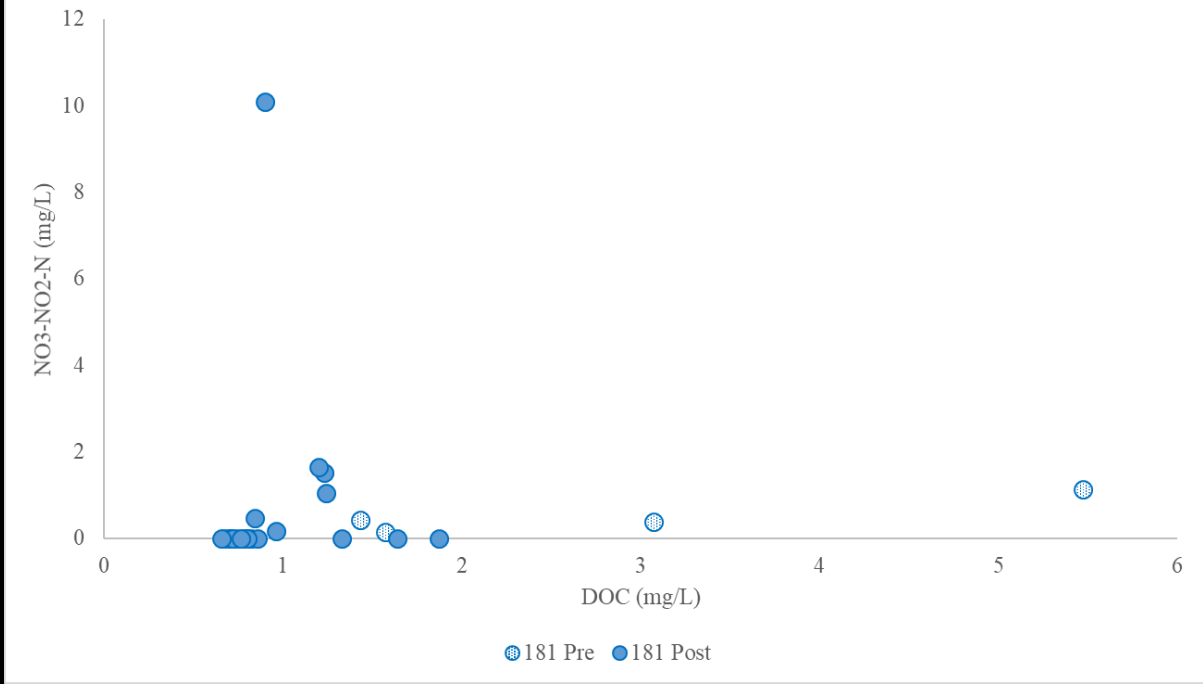
Well **199** is located upslope in the **unrestored area**, whereas **181** and **189** are inside the **restored area**.

Well **181** had a depth of 4.9 ft, but was replaced post-restoration so the depth became 2.6. The screened interval (SI) begins at 1.6 ft (relict) - The soil type is Newark.

Well **189** had a pre-restoration depth of 7.8 ft, and was replaced post-restoration but the depth remained the same. The screened SI begins at 3 ft (C horizon/bedrock) - The soil type is Newark.

Well **199** had a depth of 12.5 ft pre-restoration. This depth became 11.5 ft when the well was replaced post-restoration but the SI begins at 6.5 ft (bedrock). The overlying soil type is Conestoga.





Pre-restoration: There appears to be higher DOC, even where NO₃-NO₂-N is low. In the unrestored area, oxic conditions at **199** may inhibit denitrification.

Post-restoration: the NO₃-NO₂-N:DOC seems to stabilize at **199** and **181**, however the ratio of nitrates to DOC at **199** are still high. **189** is located directly downslope of **194** which may explain some of the trends we see along the left side of the graph.

Introduction: High-Low Residence

Groundwater residence time at these wells was longer than the catchment average pre-restoration occurred, but turnover seemed to *increase* post-restoration resulting in a lower than average residence time. Thus, the **groundwater flow paths at these wells became *more active* post-restoration.**

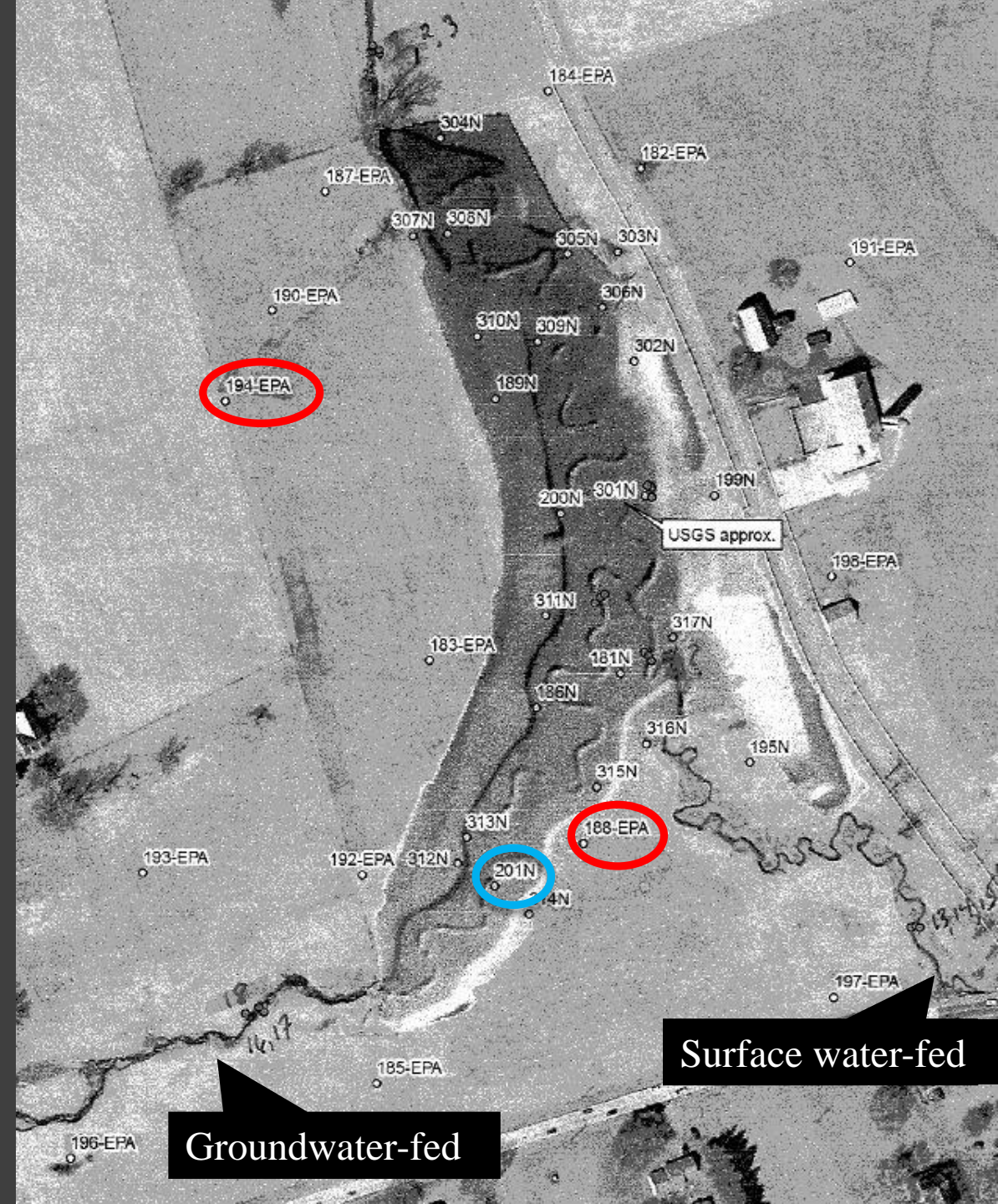
Wells **188** and **194** are located in the **unrestored** area, whereas **201** is inside the **restored** area.

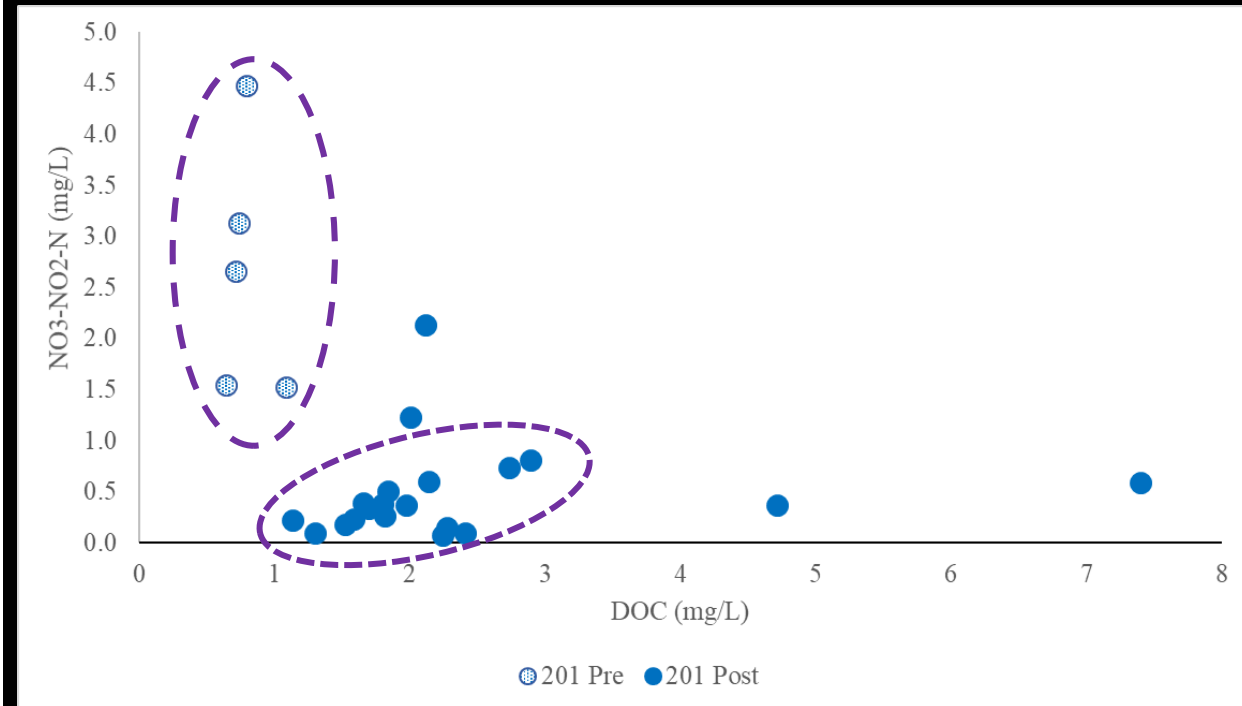
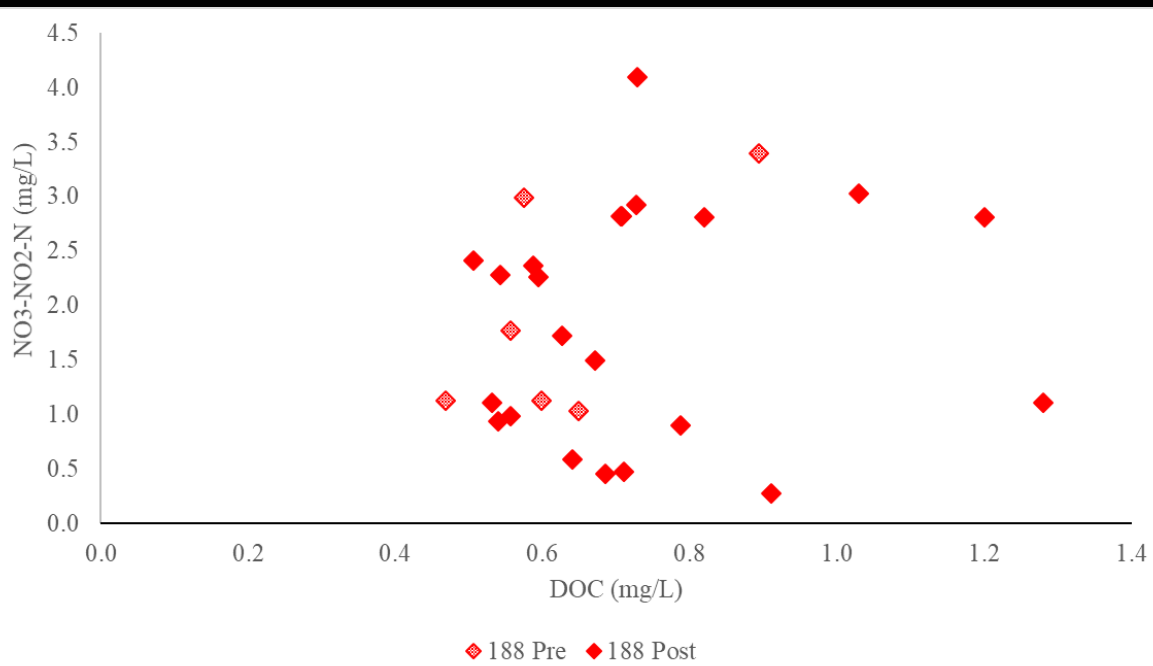
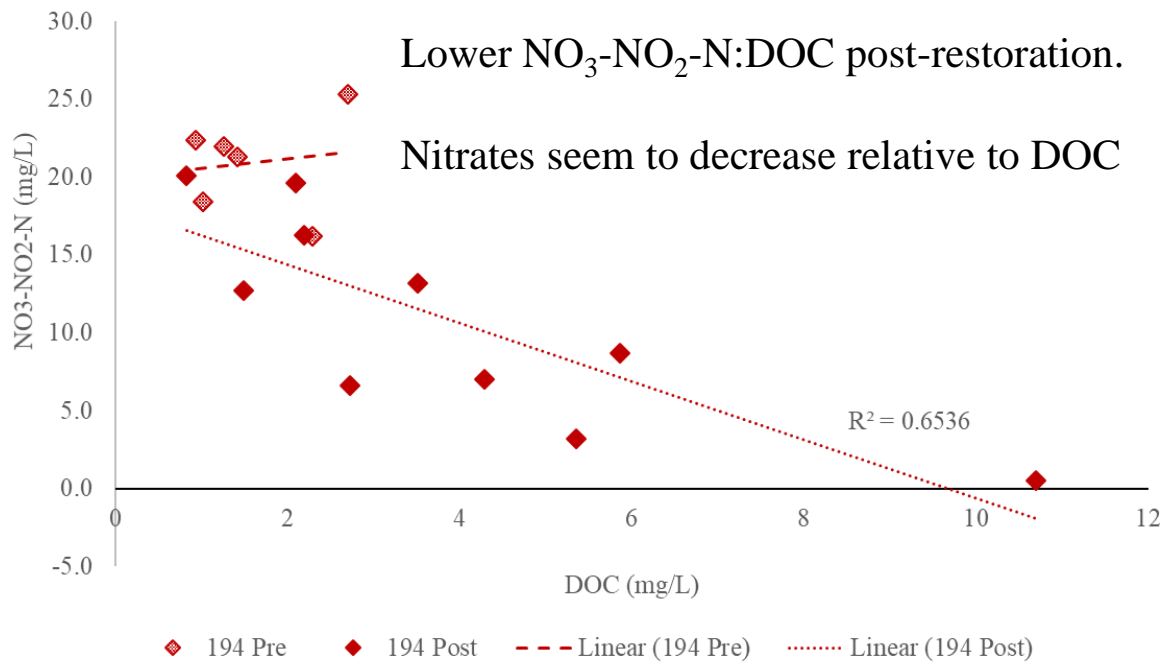
Well **194** is located further uphill of the restoration area than **188**. This spatial difference may be important as we consider the restoration zone of influence.

Well **188** has a depth of 11.7 ft, but screened interval (SI) begins at 6.7 ft (C horizon) - The soil type is Newark.

Well **194** has a depth of 14.0 ft, but the SI begins at 9 ft (bedrock). The overlying soil type is Pequea.

Well **201** depth was 11.8 ft pre-restoration and was re-drilled to 3.5 ft (SI begins at 2.5 ft) post-restoration. The soil type is Newark.





Co-located wells **188** and **201** demonstrate consistent low NO₃-NO₂-N and stable DOC relative to **194**, which is located further from the restored area

Introduction: High-High Residence

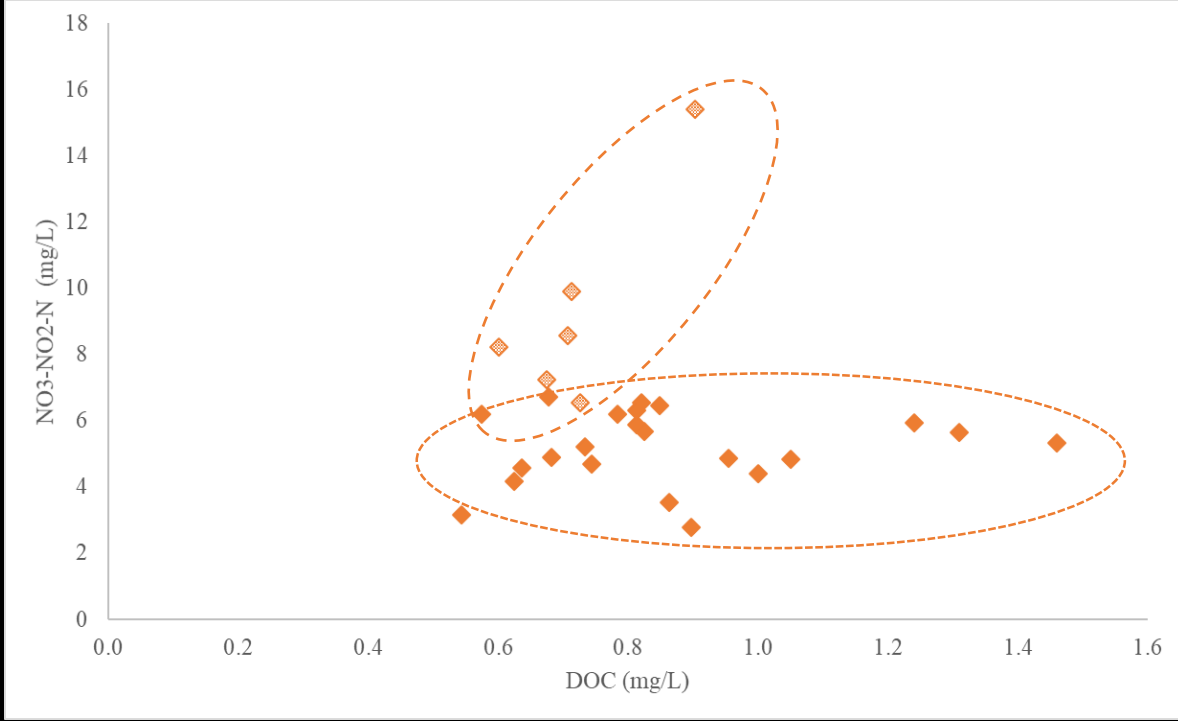
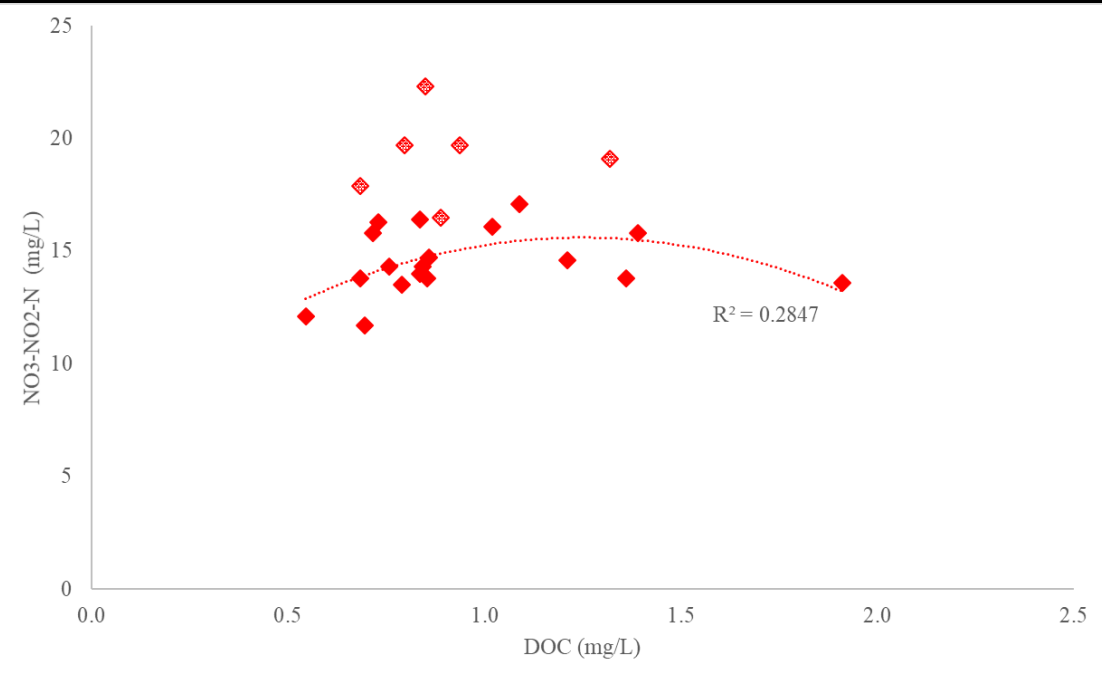
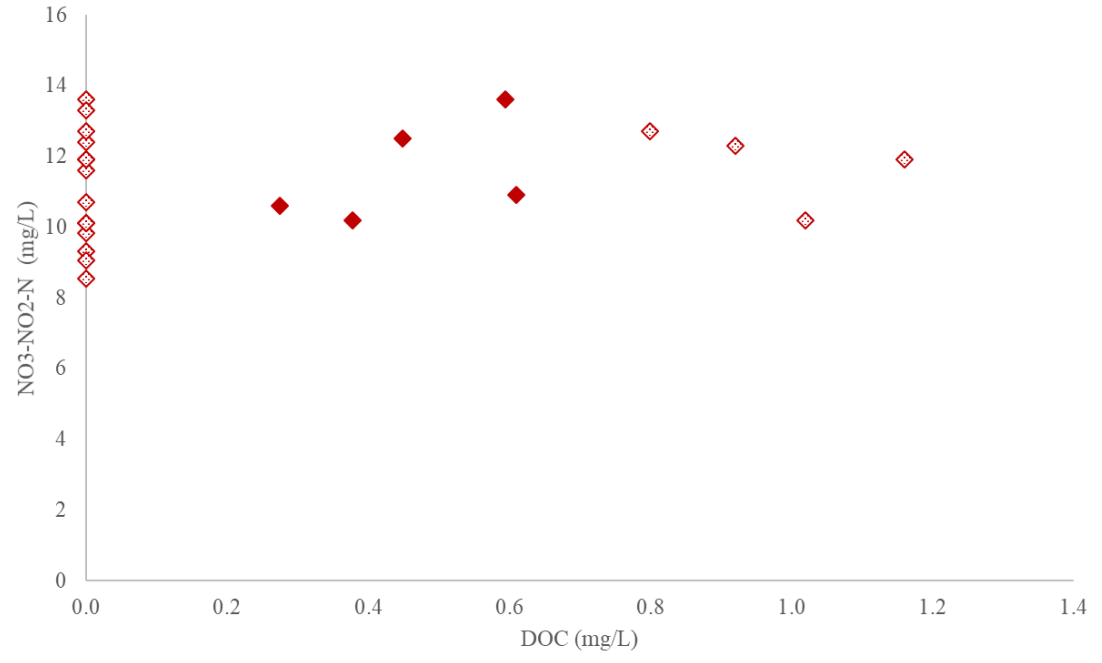
Groundwater residence time at these wells was longer than the catchment average both before and after the restoration occurred. **Groundwater at these locations assumes relatively *inactive* flow paths.**

All wells are located outside of the restored area.

Well **195** (pre depth: 10 ft, post: 8.0 ft) and well **198** (depth: 9.1 ft) are located nearest to the surface water-fed tributary (eastern branch). SI begins at 3 and 5 ft, respectively.

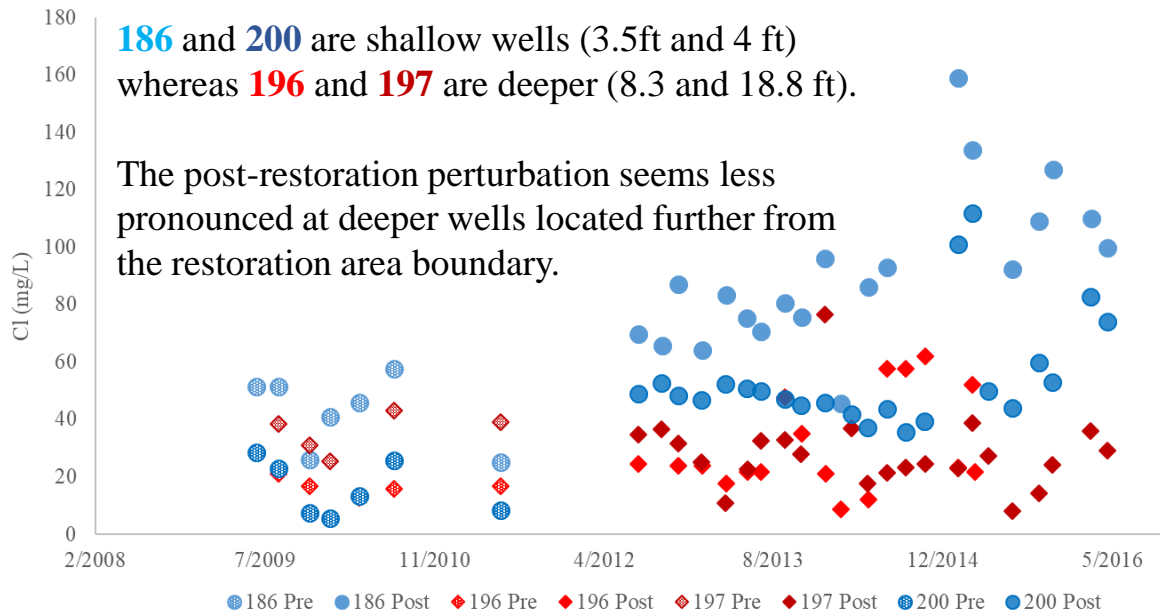
Well **185** (depth: 6.4ft, SI at 5 ft) is located nearest to the groundwater-fed tributary (western branch). This well was decommissioned in 2015.



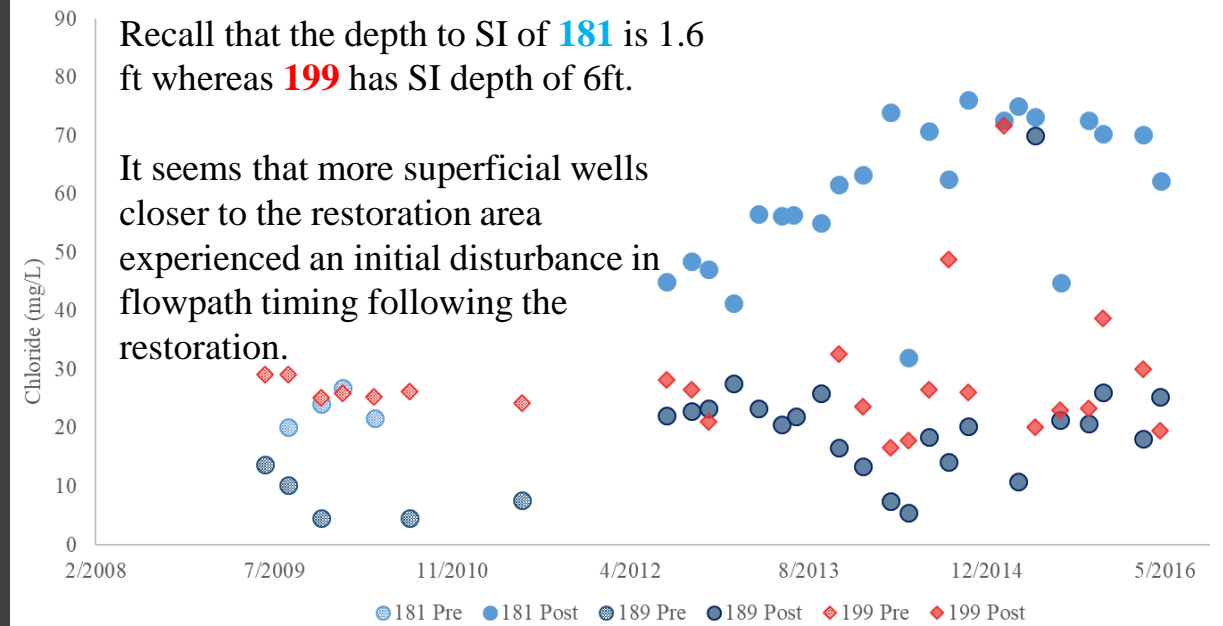


DOC is consistently low (< 2mg/L) relative to NO₃-NO₂-N at these wells located outside of the restored area with long residence times

Low - Low Residence, CI

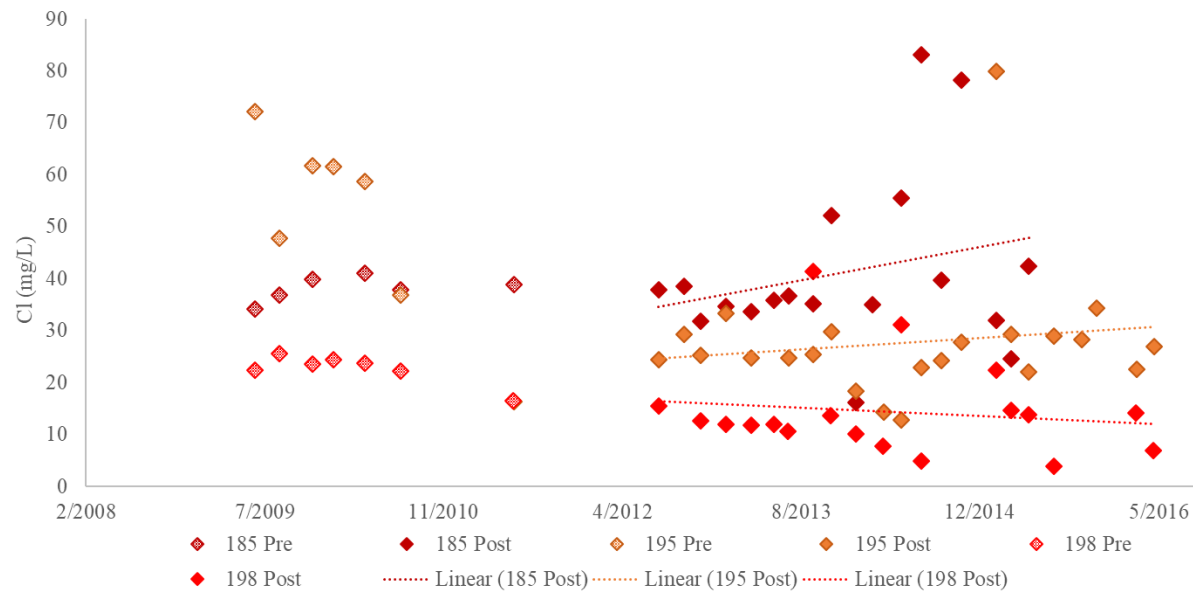


Low - High Residence, CI

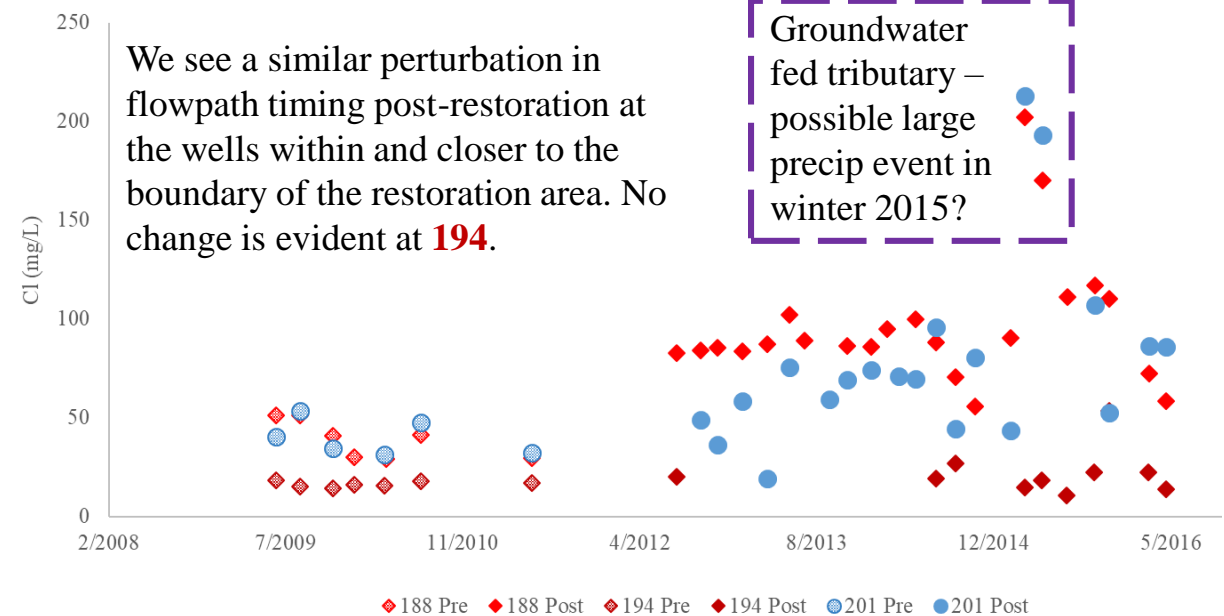


Chloride is a useful and complementary tracer to stable isotopes because it remains relatively inert in systems during recharge events.

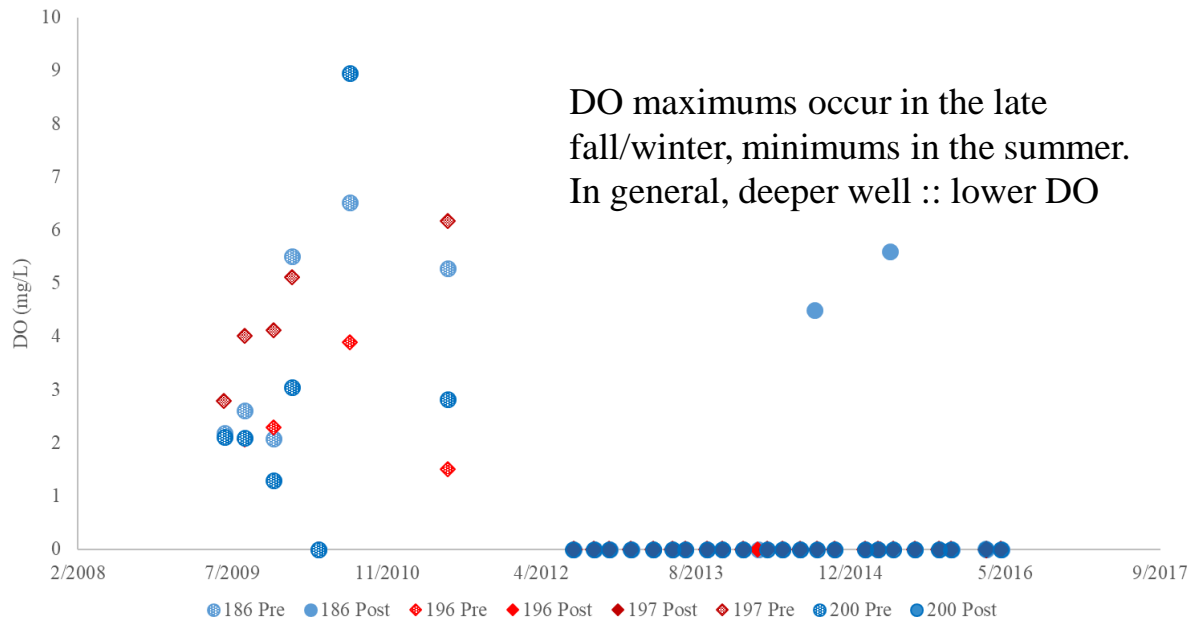
High - High Residence, CI



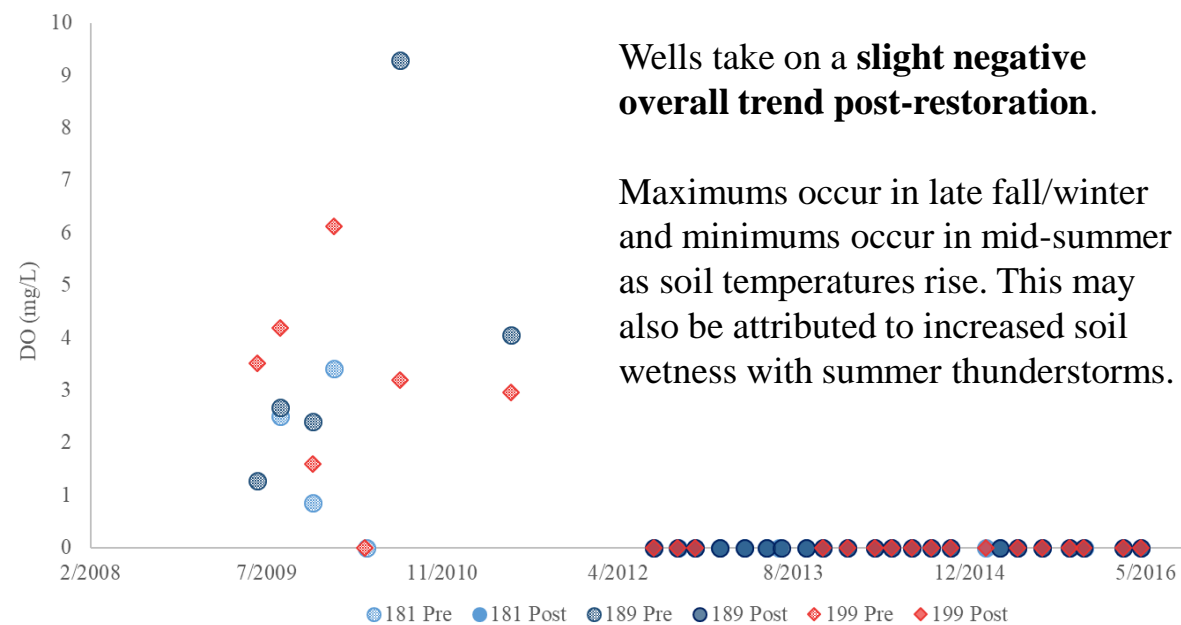
High - Low Residence, CI



Low - Low Residence, DO

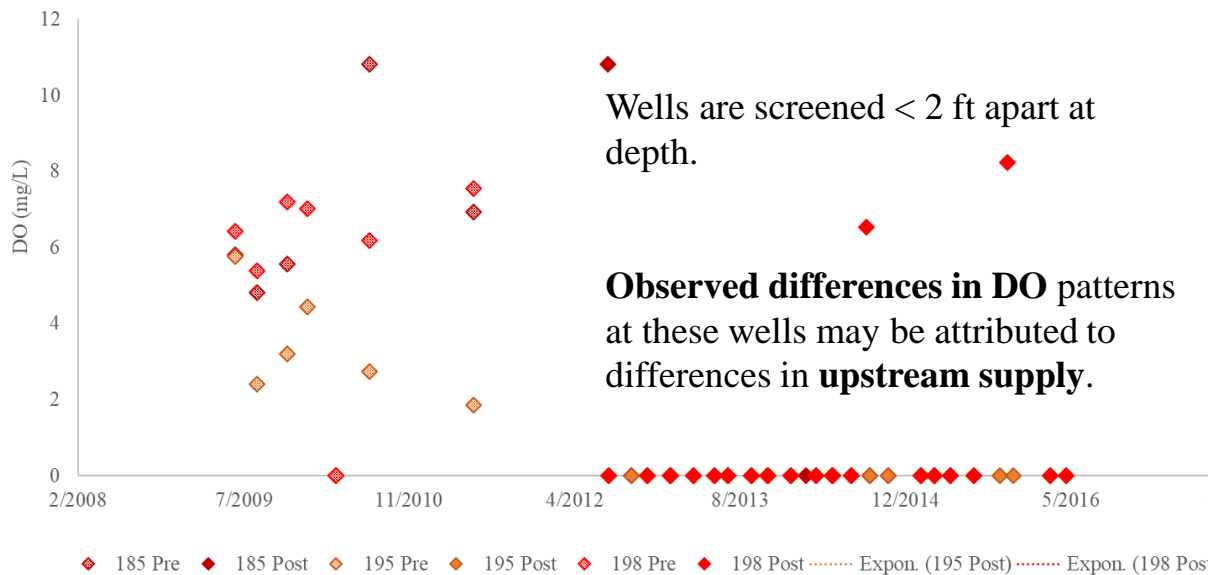


Low - High Residence, DO

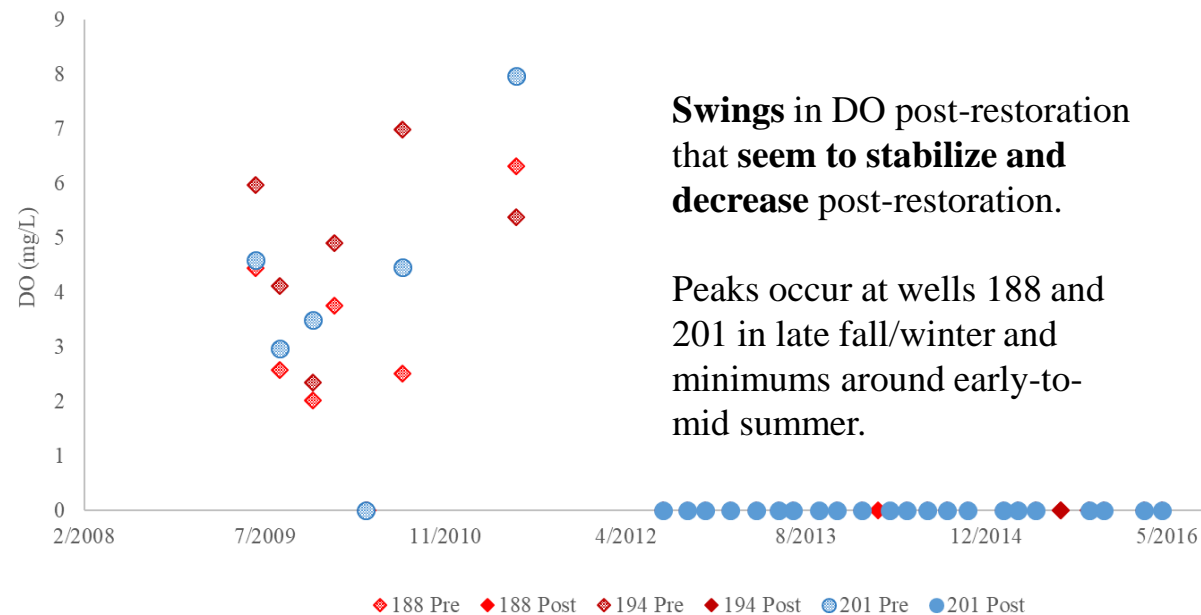


Dissolved oxygen appears to have decreased and stabilized across all groundwater wells post-restoration.

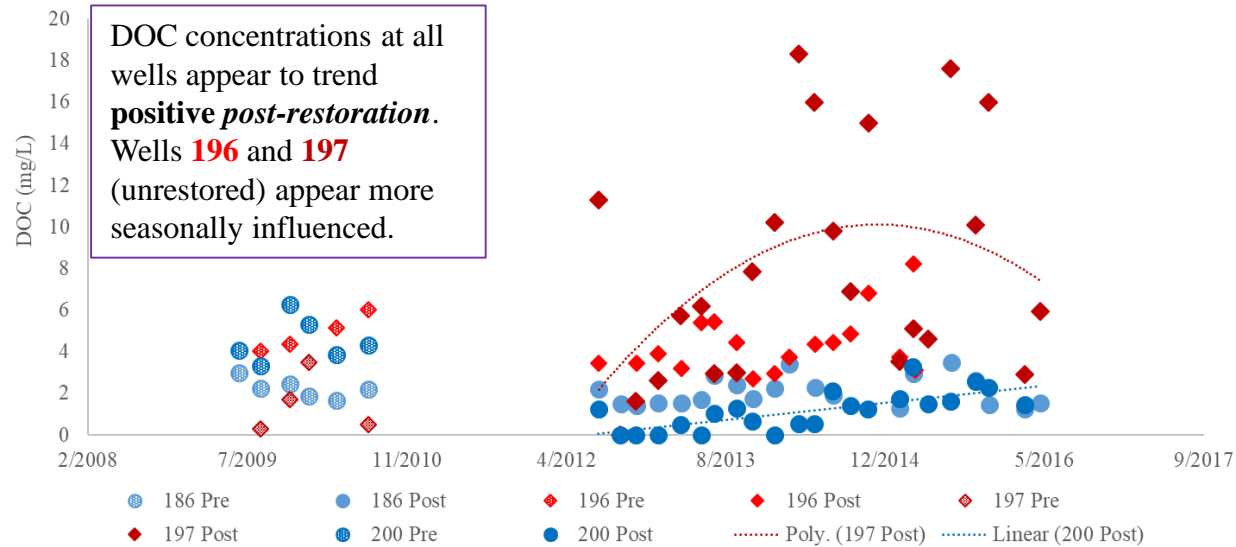
High - High Residence, DO



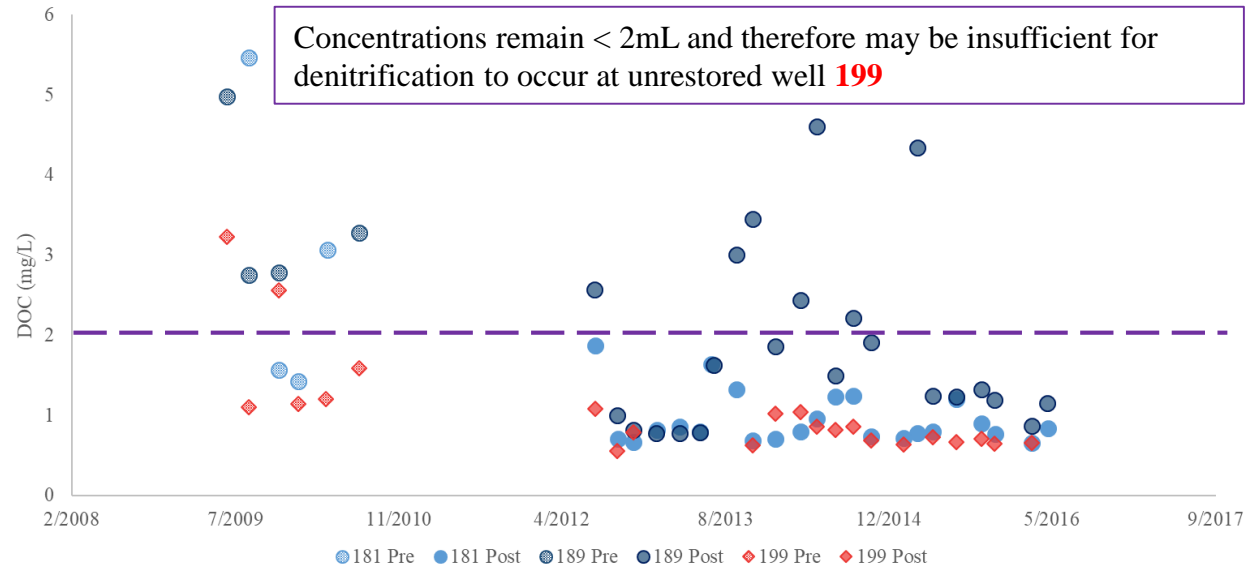
High - Low Residence, DO



Low - Low Residence, DOC

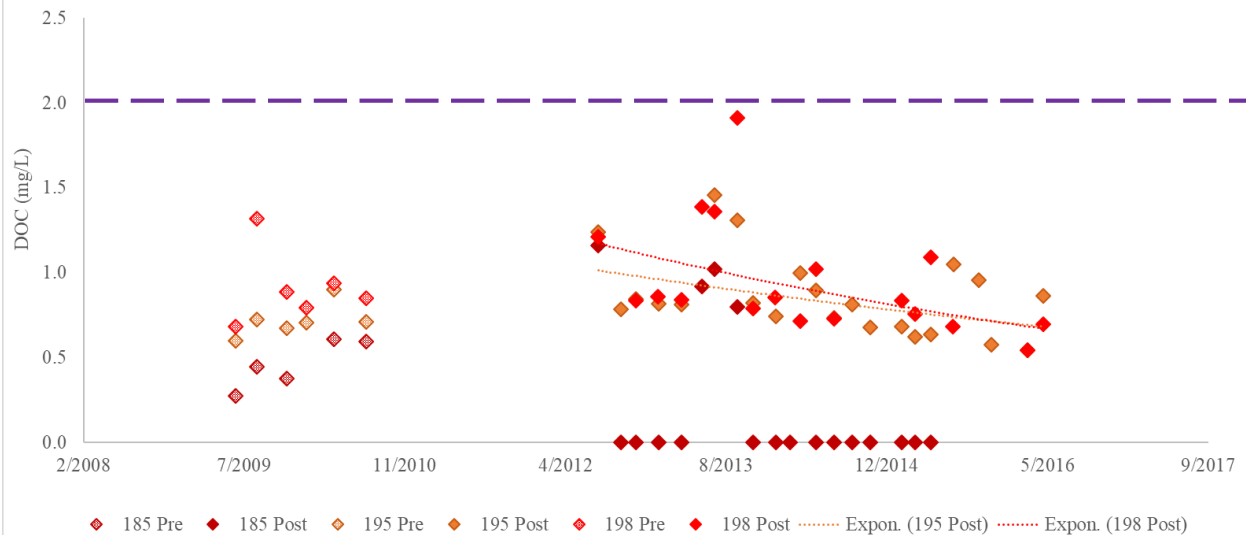


Low - High Residence, DOC

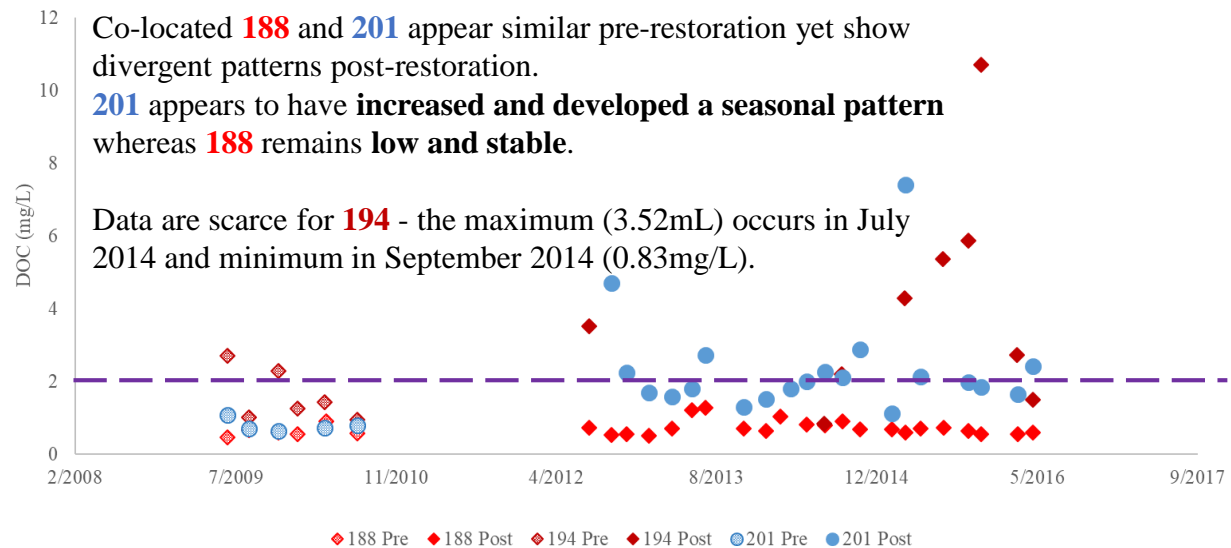


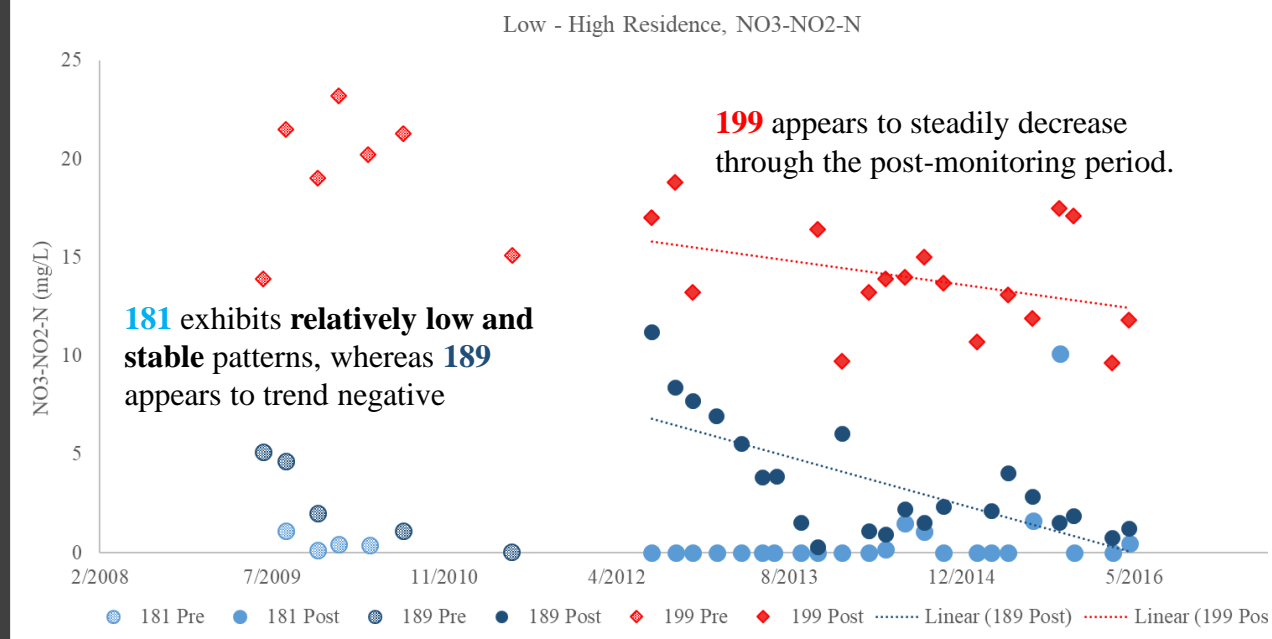
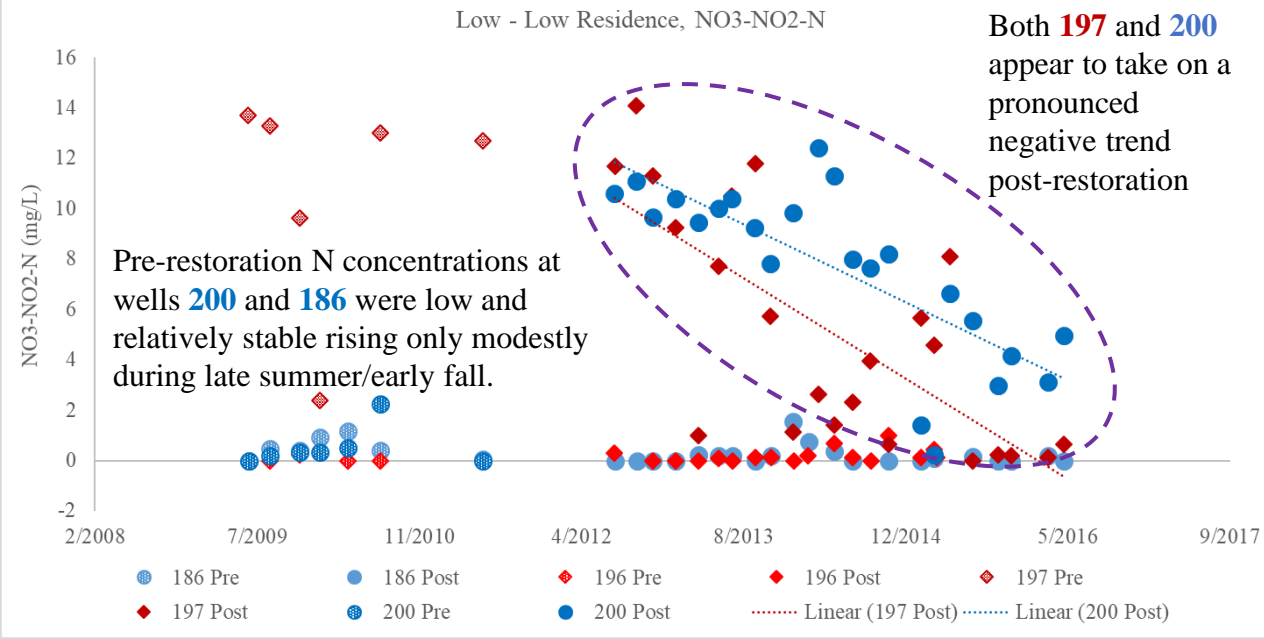
Wells appear to be carbon-limited. If groundwater remains in residence for longer periods of time, dilution of carbon inputs may play a factor in denitrification.

High - High Residence, DOC



High - Low Residence, DOC





Seasonal patterns in N are evident at some wells. Peaks occurred in late spring-to-summer and lows in winter possibly related to DO +/- DOC.

